ARTICLE IN PRESS



Available online at www.sciencedirect.com

SCIENCE () DIRECT

Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx dynamics of atmospheres and oceans

www.elsevier.com/locate/dynatmoce

Localized multiscale energy and vorticity analysis I. Fundamentals

X. San Liang^{a,*}, Allan R. Robinson^{a,b}

^a Division of Engineering and Applied Sciences, Harvard University, 29 Oxford Street, Cambridge, MA 02138, USA

^b Department of Earth and Planetary Sciences, Harvard University, Cambridge, MA, USA

Received 6 October 2003; received in revised form 10 December 2004; accepted 17 December 2004

10 Abstract

5

6

7

8

q

A new methodology, multiscale energy and vorticity analysis (MS-EVA), is developed to in-11 vestigate the inference of fundamental processes from oceanic or atmospheric data for complex 12 dynamics which are nonlinear, time and space intermittent, and involve multiscale interactions. 13 Based on a localized orthogonal complementary subspace decomposition through the multiscale 14 window transform (MWT), MS-EVA is real problem-oriented and objective in nature. The de-15 velopment begins with an introduction of the concepts of scale and scale window and the de-16 composition of variables on scale windows. We then derive the evolution equations for multi-17 scale kinetic and available potential energies and enstrophy. The phase oscillation reflected on 18 the horizontal maps from Galilean transformation is removed with a 2D large-scale window 19 synthesis. The resulting energetic terms are analyzed and interpreted. These terms, after being 20 carefully classified, provide four types of processes: transport, transfer, conversion, and dissipa-21 tion/diffusion. The key to this classification is the transfer-transport separation, which is made 22 possible by looking for a special type of transfer, the so-called *perfect transfer*. The intricate 23 energy source information involved in perfect transfers is differentiated through an interaction 24 analysis. 25

The transfer, transport, and conversion processes form the basis of dynamical interpretation for GFD problems. They redistribute energy in the phase space, physical space, and space of energy types. These processes are all referred to in a context local in space and time, and therefore can be

* Corresponding author. Tel.: +1 617 495 2899; fax: +1 617 495 5192. *E-mail address:* liang@deas.harvard.edu (X. San Liang).

^{1 0377-0265/\$ -} see front matter Elsevier B.V. All right reserved.

² doi:10.1016/j.dynatmoce.2004.12.004

2

ARTICLE IN PRESS

- easily applied to real ocean problems. When the dynamics of interest is on a global or duration scale,
- ²⁹ MS-EVA is reduced to a classical Reynolds-type energetics formalism.
- 30 Elsevier B.V. All right reserved.
- 31 Keywords: MS-EVA; Multiscale window transform; Perfect transfer; Interaction analysis
- 32

33 1. Introduction

Energy and vorticity analysis is a widely used approach in the diagnosis of geophysical 34 fluid processes. During past decades, much work has been done along this line, examples 35 including Holland and Lin (1975), Harrison and Robinson (1978), Plumb (1983), Pinardi 36 and Robinson (1986), Spall (1989), Cronin and Watts (1996), to name but a few. While these 37 classical analyses have been successful in their respective applications, real ocean processes 38 usually appear in more complex forms, involving interactions among multiple scales and 39 tending to be intermittent in space and time. In order to investigate ocean problems on a 40 generic basis, capabilities of classical energetic analyses need to be expanded to appropri-41 ately incorporate and faithfully represent all these processes. This forms the objective of 42 this work. 43

We develop a new methodology, multiscale energy and vorticity analysis (MS-EVA), 44 to fulfill this objective. MS-EVA is a generic approach for the investigation of multiscale 45 nonlinear interactive oceanic processes which occur locally in space and time. It aims 46 to explore pattern generation and energy and enstrophy budgets, and to unravel the in-47 tricate relationships among events on different scales and in different locations. In the 48 sequels to this paper (referred to as LR1), Liang and Robinson (2003a,b) (LR2 and LR3 49 hereafter), we will show how MS-EVA can be utilized for instability analysis and how 50 it can be applied to solve real ocean problems which would otherwise be difficult to 51 solve. 52

In order to be real problem-oriented, MS-EVA should contain full physics. Approximations such as linearization are thus not allowed. It must also have a multiscale representation which retains time and space localization. In other words, the representation should retain time intermittency, and should be able to handle events occurring on limited, irregular and time dependent domains. This makes MS-EVA distinctly different from classical formalism.

MS-EVA should also be scale windowed, i.e., the multiscale decomposition must be able 59 to represent events occurring coherently on scale ranges, or scale windows. Loosely speak-60 ing, a scale window is simply a subspace with a certain range of scales. A rigorous definition 61 is deferred to Section 2. In general, GFD processes tend to occur on scale windows, rather 62 than individual scales. We refer to this phenomenon as scale windowing. Scale windowing 63 requires a special bulk treatment of energy rather than individual scale representations, as 64 transfers between individual scales belonging respectively to different windows could take 65 a direction opposite to the overall transfer between these windows. 66 Multiscale events could be represented in different forms. One of the most frequently

Multiscale events could be represented in different forms. One of the most frequently used is wave representation (e.g., Fourier analysis), which transforms events onto many

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

individual scales; another frequently used form is called eddy representation (Tennekes and 69 Lumley, 1972), in which a process is decomposed into a large-scale part and an eddy part, 70 each part involving a range of scales. Because of its scale window nature, we need an eddy 71 representation for MS-EVA. The resulting energetics will be similar to those of Reynolds 72 formulation, except that the latter is in a statistical context. 73 To summarize, it is required that MS-EVA handle fairly generic processes in the sense 74 of multiscale windowing, spatial localization, and temporal intermittency; as well as re-75 tain full physics. Correspondingly an analysis tool is needed in the MS-EVA formulation 76 such that all these requirements are met. We will tackle this problem in a spirit simi-77 lar to the wavelet transform, a localized analysis which has been successfully applied to 78 studying energetics for individual scales (e.g., Iima and Toh, 1995; Fournier, 1999). Specif-79 ically, we need to generalize the wavelet analysis to handle window or eddy decomposition. 80 The challenge is how to incorporate into a window the transform coefficients (and hence 81

energies) of an orthonormal wavelet transform which are defined discretely at different 82 locations for different scales, while retaining a resolution satisfactory to the problem. (Or-83 thonormality is essential to keep energy conserved.) The next section is intended to deal 84 with this issue. The new analysis tool thus constructed will be termed *multiscale window* 85 transform, or MWT for short. The whole problem is now reduced to first the building 86 of MWT, and then the development of MS-EVA with the MWT. In Sections 3–7, we 87 apply MWT to derive the laws that govern the multiscale energy evolutions. The multi-88 scale decomposition is principally in time, but with a horizontal treatment which preserves 89 spatial localization. Time scale decomposition has been a common practice and meteo-90 rologists find it useful for clarifying atmospheric processes. We choose to do so in order 91 to make contacts with the widely used Reynolds averaging formalism, and more impor-92 tantly, to have the concept *scale* unambiguously defined (cf. Section 2.1), avoiding extra 93 assumptions such as space isotropy or anisotropy. Among these sections, Section 3 is de-94 voted to define energy on scale windows, and Section 4 is for a primary treatment with 95 the nonlinear terms. The multiscale kinetic and potential energy equations are first de-96 rived in Sections 5 and 6 based on a time decomposition, and then modified to resolve 97 the spatial issue with a horizontal synthesis (Section 7). In Section 8, we demonstrate 98 how these equations are connected to energetics in the classical formalism. This section 99 is followed by an interaction analysis for the differentiation of transfer sources (Section 100 9), which allows a description of the energetic scenario with our MS-EVA analysis in 101 both physical and phase spaces (Section 10). As "vorticity" furnishes yet another part of 102 MS-EVA, in Section 11 we briefly present how enstrophy evolves on multiple scale win-103 dows. This work is summarized in Section 12, where prospects for application are outlined 104 as well. 105

2. Multiscale window analysis and marginalization

¹⁰⁷ In this section, we introduce the concept of scale window, multiscale window transform

(MWT), and some properties of the MWT, particularly a property referred to as marginal-

¹⁰⁹ ization. A thorough and rigorous treatment is beyond the scope of this paper. For details, the reader is referred to Liang (2002) (L02 hereafter) and Liang and Anderson (2003).

4

ARTICLE IN PRESS

110 2.1. Scale and scale window

The introduction of MWT relies on how a scale is defined. In this context, our definition of *scale* is based on a modified wavelet analysis (cf., Hernández and Weiss, 1996). For convenience, we limit the initial discussion to 1D functions. The multi-dimensional case is a direct extension and can be found in L02, Section 2.7. For any function $p(t) \in L_2[0, 1]$,¹ it can been analyzed as (L02):

116
$$p(t) = \sum_{j=0}^{+\infty} \sum_{n=0}^{2^{j}\varrho-1} \tilde{p}_{n}^{j} \psi_{n}^{\varrho,j}(t), \qquad t \in [0,1],$$
(1)

117 where

118
$$\psi_n^{\varrho,j}(t) = \sum_{\ell=-\infty}^{+\infty} 2^{j/2} \psi[2^j(t+\varrho\ell) - n], \qquad n = 0, 1, \dots, 2^j \varrho - 1$$
 (2)

and ψ is some orthonormalized wavelet function.² Here we choose it to be the one built from cubic splines, which is shown in Fig. 1a. The "period" ρ has two choices only: one is $\rho = 1$, which gives a periodic extension of the signal of interest from [0, 1] to the whole real line \mathbb{R} ; another is $\rho = 2$, corresponding to an extension by reflection, which is also an "even periodization" of the finite signal to \mathbb{R} (see L02 for details).

The distribution of $\psi_n^{1,j}(t)$ with j = 2, 4, 6 is shown in Fig. 1b. Each *j* corresponds to a quantity 2^{-j} , which can be used to define a time metric to relate the passage of temporal events since a selected epoch. We call this *j* a *scale level*, and 2^{-j} the corresponding *scale* over [0, 1].

Given the scale as conceptualized, we proceed to define scale windows. In the analysis (1), we can group together those parts with a certain range of scale levels, say, $(j_1, j_1 + 1, ..., j_2)$, to form a subspace of $L_2[0, 1]$. This subspace is called a *scale window* of $L_2[0, 1]$ in L02 with scale levels ranging from j_1 to j_2 . In doing this, any function in $L_2[0, 1]$, say p(t), can be decomposed into a sum of several parts, each encompassing exclusively features on a certain window of scales. Specifically for this work, we define three scale windows:

• large-scale window: $0 \le j \le j_0$,

- meso-scale window: $j_0 < j \le j_1$,
- sub-mesoscale window: $j_1 < j \le j_2$.

The scale level bounds j_0 , j_1 , j_2 are set according to the problem under consideration. Particularly, j_2 corresponds to the finest resolution (sampling interval 2^{-j_2}) permissible by the given finite signals. By projecting p(t) onto these three windows, we obtain its large-scale, meso-scale, and sub-mesoscale features, respectively. This decomposition is orthogonal, so the total energy thus yielded is conserved.

¹ The notation $L_2[0, 1]$ is used to indicate the space of square integrable functions defined on [0, 1].

² This is to say, $\{\psi(t - \ell), \ell \in \mathbb{Z}\}$ (\mathbb{Z} the set of integers) forms an orthonormal set.





Fig. 1. Scaling and wavelet functions (a) and their corresponding periodized bases $(\varrho = 1) \{\phi_n^{\varrho,j}(t)\}_n$ (left panel) and $\{\psi_n^{\varrho,j}(t)\}_n$ (right panel) with scale levels j = 2 (top), j = 4 (middle), and j = 6 (bottom), respectively (b). The scaling and wavelet functions ϕ and ψ are constructed from cubic splines (see Liang, 2002, Section 2.5).

143 2.2. Multiscale window transform

Scale windows are defined with the aid of wavelet basis, but the definition of multiscale window transform does not follow the same line because of the difficulty we have described

DTD 5

ARTICLE IN PRES

6

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

in the introduction, i.e., that orthonormal wavelet transform coefficients are defined dis-146 cretely on different locations for different scales. To circumvent this problem, we make a 147 direct sum of the subspaces spanned by the wavelet basis $\{\psi_n^{\varrho,m}(t)\}_n$, for all $m \leq j$. The 148 shift-invariant basis of the resulting subspace can be shown to be $\phi_n^{\varrho,j}(t)$ (L02), which is 149 the periodization [cf. (2)] of some $\phi(t)$, the orthonormal scaling function in company with 150 the wavelet function $\psi(t)$. Here ϕ is an orthonormalized cubic spline, as shown in Fig. 1a. 151 We utilize the $\phi_n^{\varrho,j}$ thus formed to fulfill our task. In the following only the related formulas 152 and equations are presented. The details are referred to L02. 153

Let V_{ϱ, j_2} indicate the total (direct sum, to be strict) of the three scale windows. It has been established by L02 that any time signal from a given GFD dataset is justifiably belonging to V_{ϱ, j_2} , with some finite level j_2 . Suppose we have $p(t) \in V_{\varrho, j_2}$. Write

157
$$\hat{p}_n^j = \int_0^{\varrho} p(t)\phi_n^{\varrho,j}(t) \,\mathrm{d}t, \quad \text{for all } 0 \le j \le j_2, \quad n = 0, 1, \dots, 2^j \varrho - 1.$$
 (3)

Given window bounds j_0, j_1, j_2 , and $p \in V_{\varrho, j_2}$, three functions can be accordingly defined:

159
$$p^{\sim 0}(t) = \sum_{n=0}^{2^{j_0}\varrho - 1} \hat{p}_n^{j_0} \phi_n^{\varrho, j_0}(t),$$
 (4)

160
$$p^{\sim 1}(t) = \sum_{n=0}^{2^{j_1} \varrho - 1} \hat{p}_n^{j_1} \phi_n^{\varrho, j_1}(t) - p^{\sim 0}(t),$$
 (5)

$$p^{\sim 2}(t) = p(t) - \sum_{n=0}^{2^{j_1} \varrho - 1} \hat{p}_n^{j_1} \phi_n^{\varrho, j_1}(t),$$
(6)

on the basis of which we will build the MWT later. As a scaling transform coefficient, \hat{p}_n^J contains all the information with scale level lower than or equal to *j*. The functions $p^{\sim 0}(t)$, $p^{\sim 1}(t)$, $p^{\sim 2}(t)$ thus defined hence include only features of p(t) on ranges $0 - j_0$, $j_0 - j_1$, and $j_1 - j_2$, respectively. For this reason, we term these functions as large-scale, meso-scale, and sub-mesoscale syntheses or reconstructions of p(t), with the notation ~ 0 , ~ 1 , and ~ 2 in the superscripts signify the corresponding large-scale, meso-scale, and sub-mesoscale windows, respectively.

¹⁶⁹ Using the multiscale window synthesis, we proceed to define a transform

170
$$\hat{p}_n^{\sim\varpi} = \int_0^{\varrho} p^{\sim\varpi}(t)\phi_n^{\varrho,j_2}(t)\,\mathrm{d}t \tag{7}$$

for windows $\varpi = 0, 1, 2, n = 0, 1, \dots, 2^{j_2 \varrho} - 1$. This is the *multiscale window transform*, or MWT for short, that we want to build. Notice here we use a periodized scaling basis at j_2 , the highest level that can be attained for a given time series. As a result, the transform coefficients have a maximal resolution in the sampled *t* direction.

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

In terms of $\hat{p}_n^{\sim \overline{\omega}}$, Eqs. (4)–(6) can be simplified as

$$p^{\sim \varpi}(t) = \sum_{n=0}^{2^{j_2} \varrho - 1} \hat{p}_n^{\sim \varpi} \phi_n^{\varrho, j_2}(t),$$
(8)

for $\overline{\omega} = 0, 1, 2$. Eqs. (7) and (8) are the transform-reconstruction pair for our MWT. For any $p \in V_{\rho, j_2}$, it can be now represented as

179
$$p(t) = \sum_{\varpi=0}^{2} \sum_{n=0}^{2^{j_2} \varrho - 1} \hat{p}_n^{\sim \varpi} \phi_n^{\varrho, j_2}(t).$$
(9)

¹⁸⁰ A final remark on the choice of extension scheme, or the "period" ρ in the analysis. In ¹⁸¹ general, we always adopt the extension by reflection $\rho = 2$, which has proved to be very ¹⁸² satisfactory. (Fig. 4 shows such an example.) If the signals given are periodic, then the ¹⁸³ periodic extension is the exact one, and hence ρ should be chosen to be 1. In case of linking ¹⁸⁴ to the classical energetic formalism, $\rho = 1$ is also usually used.

185 2.3. MWT properties and marginalization

Multiscale window transform has many properties. In the following we present two of them which will be used later in the MS-EVA development (for proofs, refer to L02).

Property 1. For any $p \in V_{\rho, j_2}$, if $j_0 = 0$, and $\rho = 1$ (periodic extension adopted), then

$$\hat{p}_n^{\sim 0} = 2^{-j_2/2} p^{\sim 0}(t) = 2^{-j_2/2} \bar{p} = constant, \quad for \ all \ n, \ and \ t, \tag{10}$$

¹⁹⁰ where the overbar stands for averaging over the duration.

¹⁹¹ **Property 2.** For p and q in V_{ϱ, j_2} ,

192
$$\mathcal{M}_n \hat{p}_n^{\sim \varpi} \hat{q}_n^{\sim \varpi} = \overline{p^{\sim \varpi}(t)q^{\sim \varpi}(t)}, \tag{11}$$

193 where

1

$$\mathcal{M}_{n}(\hat{p}_{n}^{\sim\varpi}\hat{q}_{n}^{\sim\varpi}) = \sum_{n=1}^{N-1} \hat{p}_{n}^{\sim\varpi}\hat{q}_{n}^{\sim\varpi} + \frac{1}{2}[\hat{p}_{0}^{\sim\varpi}\hat{q}_{0}^{\sim\varpi} + \hat{p}_{N}^{\sim\varpi}\hat{q}_{N}^{\sim\varpi}]. \quad (N = 2^{j_{2}}) \quad (12)$$

Property 1 states that when $j_0 = 0$ and a periodic extension is used, the large-scale 195 window synthesis is simply the duration average. Property 2 involves a special summation 196 over [0, N] (corresponding to $t \in [0, 1]$), which we will call *marginalization* hereafter. 197 The word "marginal" has been used in literature to describe the overall feature of a 198 localized transform (e.g., Huang et al., 1999). We extend this convention to establish an 199 easy reference for the operator \mathcal{M}_n . Property 2 can now be restated as: a product of two 200 multiscale window transforms followed by a marginalization is equal to the product of 201 their corresponding syntheses averaged over the duration. For convenience, this property 202 will be referred to as property of marginalization. 203

We close this section by making a comparison between our MWT and wavelet analysis. The commonality is, of course, that both of them are localized on the definition

8

228

2

2

231

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

domain. The first and largest difference between them is that the MWT is not a trans-206 form in the usual sense. It is an orthogonal complementary subspace decomposition, and 207 as a result, the MWT coefficients contain information for a range of scales, instead of 208 a single scale. For this reason, it is required that three scale bounds be specified a pri-209 ori in constructing the windows. A useful way to do this is through wavelet spectrum 210 analysis, as is used in LR3. Secondly, the MWT transform is projected on V_{ρ, j_2} , so trans-211 form coefficients obtained for all the windows have the same resolution-the maximal 212 resolution allowed for the signal. This is in contrast to wavelet analysis, whose transform 213 coefficients have different resolution on different scales. We will see soon that, this maxi-214 mized resolution in MWT transform coefficients puts the embedded phase oscillation under 215 control. 216

217 **3. Multiscale energies**

Beginning this section through Section 7, we will derive the equations that gov-218 ern the multiscale energy evolutions. The whole formulation is principally based on 219 a time decomposition, but with an appropriate filtering in the horizontal dimensions. 220 It involves a definition of energies on different scale windows, a classification of dis-221 tinct processes from the nonlinear convective terms, a derivation of time windowed 222 energetic equations, and a horizontal treatment of these equations with a space win-223 dow reconstruction. In this section, we define the energies for the three time scale 224 windows. 225

226 3.1. Primitive equations and kinetic and available potential energies

²²⁷ The governing equations adopted in this study are:

$$\frac{\partial \underline{\mathbf{v}}}{\partial t} = -\nabla \cdot (\underline{\mathbf{v}} \,\underline{\mathbf{v}}) - \frac{\partial (w \,\underline{\mathbf{v}})}{\partial z} - f \underline{\mathbf{k}} \wedge \underline{\mathbf{v}} - \frac{1}{\rho_0} \nabla P + \underline{\mathbf{F}}_{mz} + \underline{\mathbf{F}}_{mh},\tag{13}$$

$$0 = \nabla \cdot \underline{\mathbf{v}} + \frac{\partial w}{\partial z},\tag{14}$$

$$0 = -\frac{\partial P}{\partial z} - \rho g, \qquad (15)$$

$$\frac{\partial \rho}{\partial t} = -\nabla \cdot (\underline{\mathbf{v}}\rho) - \frac{\partial (w\rho)}{\partial z} + \frac{N^2 \rho_0}{g} w + F_{\rho z} + F_{\rho h}, \qquad (16)$$

where $\underline{\mathbf{v}} = (u, v)$ is the horizontal velocity vector, $\nabla = \underline{\mathbf{i}} \frac{\partial}{\partial x} + \underline{\mathbf{j}} \frac{\partial}{\partial y}$ the horizontal gradient operator, $N = \left(-\frac{g}{\rho_0} \frac{\partial \bar{\rho}}{\partial z}\right)^{1/2}$ the buoyancy frequency ($\bar{\rho} = \bar{\rho}(z)$ is the stationary density profile), ρ the density perturbation with $\bar{\rho}$ excluded, and *P* the dynamic pressure. All the other notations are conventional. The friction and diffusion terms are just symbolically expressed. The treatment of these subgrid processes in a multiscale setting is not considered in this paper. From Eqs. (13) and (14), it is easy to obtain the equations that govern the evolution

260

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

of two quadratic quantities: $K = \frac{1}{2} \mathbf{v} \cdot \mathbf{v}$, and $A = \frac{1}{2} \frac{g^2}{\rho_0^2 N^2} \rho^2$ (see Spall, 1989). These are the total kinetic energy (KE) and available potential energy (APE), given the location in space and time. The essence of this study is to investigate how KE and APE are distributed simultaneously in the physical and phase spaces.

242 3.2. Multiscale energies

Multiscale window transforms equipped with the marginalization property (11) allow a simple representation of energy for each scale window $\overline{\omega} = 0, 1, 2$. For a scalar field $S(t) \in V_{\varrho, j_2}$, let $E_n^{\overline{\omega}*} = (\hat{S}_n^{-\overline{\omega}})^2$. By (11),

$$\mathcal{M}_n E_n^{\varpi *} = \int_0^1 \left[S^{\sim \varpi}(t) \right]^2 \mathrm{d}t, \tag{17}$$

which is essentially the energy of *S* on window ϖ (up to some constant factor) integrated with respect to *t* over [0, 1). Recall \mathcal{M}_n is a special sum over the 2^{j_2} discrete equi-distance locations $n = 0, 1, \ldots, 2^{j_2} - 1$. $E_n^{\varpi*}$ thus can be viewed as the energy on window ϖ summarized over a small interval of length $\Delta t = 2^{-j_2}$ around location $t = 2^{-j_2}n$. An energy variable for window ϖ at time $2^{-j_2}n$ consistent with the fields at that location is therefore a locally averaged quantity

253
$$E_n^{\varpi} = \frac{1}{\Delta t} E_n^{\varpi *} = 2^{j_2} \cdot (\hat{S}_n^{\sim \varpi})^2,$$
 (18)

for all $\varpi = 0, 1, 2$. It is easy to establish that

255
$$\mathcal{M}_n(E_n^0 + E_n^1 + E_n^2)\Delta t = \int_0^1 S^2(t) \,\mathrm{d}t.$$
(19)

²⁵⁶ This is to say, the energy thus defined is conserved.

In the same spirit, the multiscale kinetic and available potential energies now can be defined as follows:

$$K_{n}^{\varpi} = \frac{1}{2} [2^{j_{2}} (\hat{u}_{n}^{\sim \varpi})^{2} + 2^{j_{2}} (\hat{v}_{n}^{\sim \varpi})^{2}] = 2^{j_{2}} \left[\frac{1}{2} \hat{\mathbf{y}}_{n}^{\sim \varpi} \cdot \frac{1}{2} \hat{\mathbf{y}}_{n}^{\sim \varpi} \right]$$
(20)

$$A_{n}^{\varpi} = 2^{j_{2}} \left[\frac{1}{2} \frac{g^{2}}{\rho_{0}^{2} N^{2}} \hat{\rho}_{n}^{\sim \varpi} \cdot \hat{\rho}_{n}^{\sim \varpi} \right] = 2^{j_{2}} \left[\frac{1}{2} c \hat{\rho}_{n}^{\sim \varpi} \hat{\rho}_{n}^{\sim \varpi} \right],$$
(21)

where the shorthand $c \equiv g^2/(\rho_0^2 N^2)$ is introduced to avoid otherwise cumbersome derivation of the potential energy equation. (Note *c* is *z*-dependent.) The purpose of the following sections are to derive the evolution laws for $K_n^{\overline{\omega}}$ and $A_n^{\overline{\omega}}$. Note the factor 2^{j_2} , which is a constant once a signal is given, provides no information essential to our dynamics analysis. In the MS-EVA derivation, we will drop it in order to avoid otherwise awkward expressions. Therefore, all the energetic terms hereafter, unless otherwise indicated, should be multiplied by 2^{j_2} before physically interpreted.

ARTICLE IN PRESS

10 X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

4. Perfect transfer and transfer–transport separation

The MS-EVA is principally developed for time, but with a horizontal treatment for spatial oscillations. Localized energetic study with a time decomposition (and the statistical formulation) raises an issue: the separation of transport from the nonlinear term-related energetics. Here by transport we mean a process which can be represented by some quantity in a form of divergence. It vanishes if integrated over a closed domain. The separation of transport is very important, since it allows the cross-scale energy transfer to come upfront.

Transfer-transport separation is not a problem in a space decomposition-based energetic formulation, e.g., the Fourier formulation. In that case the analysis over the space has already eliminated the transport, and as a result, the summation of the triad interaction terms over all the possible scales vanishes. This problem surfaces in a localized time-based formulation when uniqueness is concerned. In this section, we will show how it is resolved.

We begin by introducing a concept, *perfect transfer process*, for our purpose. The so-279 called *perfect transfer* is a family of multiscale energetic terms which vanish upon sum-280 mation over all the scale windows and marginalization over the sampled time locations. A 281 perfect transfer process, or simply perfect transfer when no confusion arises in the context, 282 is then a process represented by perfect transfer term(s). Perfect transfers move energy from 283 window to window without destroying or generating energy as a whole. They represent a 284 kind of redistribution process among multiple scale windows. In terms of physical signifi-285 cance, the concept of perfect transfer is a natural choice. We are thence motivated to seek 286 through a larger class of "transfer processes" for perfect transfers, which set a constraint 287 for transport-transfer separation and hence help to solve the above uniqueness problem. 288

For a detailed derivation of the transport-transfer separation, refer to Liang and Robinson (2003c). Briefly cited here is the result with some modification to the needs in our context. The idea is that, for an incompressible fluid flow, we can have the nonlinear-term related energetics separated into a transport plus a perfect transfer, and the separation is unique. For simplicity, consider a scalar field S = S(t, x, y). Suppose it is simply advected by an incompressible 2D flow \underline{v} , i.e., the evolution is governed by

²⁹⁵
$$\frac{\partial S}{\partial t} = -\nabla \cdot (\underline{\mathbf{v}}S), \quad \nabla \cdot \underline{\mathbf{v}} = 0.$$
 (22)

Let $E_n^{\varpi} = \frac{1}{2} (\hat{S}_n^{\sim \varpi})^2$ be its energy (variance) at time location *n* on scale window ϖ . The evolution of E_n^{ϖ} can be easily obtained by making a transform of the equation followed by a product with $\hat{S}_n^{\sim \varpi}$. We are tasked to separate the resulting triple product term

NL =
$$-\hat{S}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}S})_n^{\sim \varpi}$$

as needed. By L02, this is done by performing the separation as

$$NL = -\nabla \cdot \underline{\mathbf{Q}}_{S_n^{\varpi}} + \left[-\hat{S}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}} \widehat{S})_n^{\sim \varpi} + \nabla \cdot \underline{\mathbf{Q}}_{S_n^{\varpi}}\right] \equiv \Delta_h Q_{S_n^{\varpi}} + T_{S_n^{\varpi}}, \tag{23}$$

302 where

303 $\underline{\mathbf{Q}}_{S_n^{\varpi}} = \lambda_c \hat{S}_n^{\sim \varpi} (\underline{\widehat{\mathbf{v}}} \widehat{S})_n^{\sim \varpi}, \qquad \lambda_c = \frac{1}{2},$ (24)

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

304 and

$$\Delta_h Q_{S_n^{\varpi}} \equiv -\nabla \cdot \mathbf{Q}_{S_n^{\varpi}} \tag{25}$$

305

$$T_{S_n^{\varpi}} \equiv -\hat{S}_n^{\sim \varpi} \nabla \cdot (\widehat{\underline{\mathbf{v}}S})_n^{\sim \varpi} + \nabla \cdot \underline{\mathbf{Q}}_{S_n^{\varpi}}.$$
(26)

307 It is easy to verify that

$$\sum_{\varpi} \mathcal{M}_n T_{S_n^{\varpi}} = 0, \qquad (27)$$

which implies that $T_{S_n^{w}}$ represents a perfect transfer process.

Eq. (23) is the transport-transfer separation for the scalar variance evolution in a 2D flow. For the 3D case, the separation is in the same form. One just needs to change the vectors and the gradient operator in (23) into their corresponding 3D counterparts.

5. Multiscale kinetic energy equation

The formulation of multiscale energetics generally follows from the derivation for the 314 evolutions of K and A. The difference lies in that here we consider our problem in the 315 phase space. Since the basis function $\phi^{\varrho,j}$, for any $0 \le j \le j_2$, is time dependent, and the 316 derivative of $\phi^{\varrho,j}$ does not in general form an orthogonal pair with $\phi^{\varrho,j}$ itself, the local time 317 change terms in the primitive equations need to be pre-treated specially before the energy 318 equations can be formulated. Similar problems also exist in Harrison and Robinson (1978)'s 319 formalism. Appearing on the left hand side of their kinetic energy equation is $\bar{\mathbf{y}} \cdot \frac{\partial \bar{\mathbf{y}}}{\partial r}$, not in 320 a form of time change of $\frac{1}{2} \mathbf{\bar{v}} \cdot \mathbf{\bar{v}}$. 321

To start, first consider $\partial \underline{v}/\partial t$. Recall that our objective is to develop a diagnostic tool for an existing dataset. Thus every differential term has to be replaced eventually by its difference counterpart. That is to say, we actually do not need to deal with $\partial \underline{v}/\partial t$ itself. Rather, it is the discretized form (space-dependence suppressed for clarity)

$$\frac{\underline{\mathbf{v}}(t+\Delta t)-\underline{\mathbf{v}}(t-\Delta t)}{2\Delta t}\equiv\delta_t\underline{\mathbf{v}}$$

•0

co

that we should pay attention to (Δt is the time step size). Viewed as functions of t, $\underline{\mathbf{v}}(t + \Delta t)$ and $\underline{\mathbf{v}}(t - \Delta t)$ make two different series and may be transformed separately. Let

326

$$\int_{0}^{\varrho} \underline{\mathbf{v}}^{\sim \varpi}(t + \Delta t) \phi_{n}^{\varrho, j_{2}}(t) \, \mathrm{d}t \equiv \underline{\hat{\mathbf{v}}}_{n+}^{\sim \varpi}, \tag{28}$$

330

$$\int_{0}^{\infty} \underline{\mathbf{v}}^{\sim \overline{\omega}} (t - \Delta t) \phi_{n}^{\varrho, j_{2}}(t) \, \mathrm{d}t \equiv \underline{\mathbf{\hat{v}}}_{n-}^{\sim \overline{\omega}}, \tag{29}$$

where ρ is the periodicity of extension ($\rho = 1$ and 2 for extensions by periodization and reflection, respectively), and define an operator $\hat{\delta}_n$ such that

$$\hat{\delta}_n \hat{\mathbf{v}}_n^{\sim \varpi} = \frac{\hat{\mathbf{v}}_{n+}^{\sim \varpi} - \hat{\mathbf{v}}_{n-}^{\sim \varpi}}{2\Delta t}.$$
(30)

DYNAT 708 1-36

12

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

 $\hat{\delta}_n \hat{\underline{v}}_n^{\infty}$ is actually the transform of $\delta_t \underline{v}$, or the rate of change of $\hat{\underline{v}}_n^{\infty}$ on its corresponding scale window. Similarly, define difference operators of the second order as follows:

$$\delta_{t^{2}} \underline{\mathbf{v}} \equiv \frac{\underline{\mathbf{v}}(t + \Delta t) - 2\underline{\mathbf{v}}(t) + \underline{\mathbf{v}}(t - \Delta t)}{(\Delta t)^{2}},$$
(31)

337

33

$$\hat{\delta}_{n^2}^2 \underline{\hat{\mathbf{y}}}_n^{\sim \overline{\omega}} \equiv \int_0^{\varrho} \delta_{t^2}^2 \underline{\mathbf{y}}^{\sim \overline{\omega}} \quad \phi_n^{\varrho, j_2}(t) \,\mathrm{d}t.$$
(32)

Now take the dot product of $\hat{\mathbf{v}}_n^{\sim \varpi}$ with $\hat{\delta}_n \hat{\mathbf{v}}_n^{\sim \varpi}$,

$$\hat{\mathbf{y}}_{n}^{\sim\varpi} \cdot \hat{\delta}_{n} \hat{\mathbf{y}}_{n}^{\sim\varpi} = \left(-\frac{\hat{\mathbf{y}}_{n+}^{\sim\varpi} - 2\hat{\mathbf{y}}_{n}^{\sim\varpi} + \hat{\mathbf{y}}_{n-}^{\infty}}{2} + \frac{\hat{\mathbf{y}}_{n+}^{\sim\varpi} + \hat{\mathbf{y}}_{n-}^{\infty}}{2} \right) \cdot \frac{\hat{\mathbf{y}}_{n+}^{\sim\varpi} - \hat{\mathbf{y}}_{n-}^{\infty}}{2\Delta t} \\
= \frac{1}{2\Delta t} \left(\frac{1}{2} \hat{\mathbf{y}}_{n+}^{\sim\varpi} \cdot \hat{\mathbf{y}}_{n+}^{\sim\varpi} - \frac{1}{2} \hat{\mathbf{y}}_{n-}^{\infty\varpi} \cdot \hat{\mathbf{y}}_{n-}^{\infty} \right) - (\Delta t)^{2} (\hat{\delta}_{n2}^{2} \hat{\mathbf{y}}_{n}^{\infty\varpi} \cdot \hat{\delta}_{n} \hat{\mathbf{y}}_{n}^{\infty\varpi}) \\
= \hat{\delta}_{n} K_{n}^{\varpi} - (\Delta t)^{2} (\hat{\delta}_{n2}^{2} \hat{\mathbf{y}}_{n}^{\infty\varpi} \cdot \hat{\delta}_{n} \hat{\mathbf{y}}_{n}^{\infty\varpi}),$$
(33)

338 where

 $K_n^{\overline{\omega}} = \frac{1}{2} \underline{\hat{\mathbf{y}}}_n^{\sim \overline{\omega}} \cdot \underline{\hat{\mathbf{y}}}_n^{\sim \overline{\omega}}$ (34)

is the kinetic energy at location *n* (in the phase space) for the window ϖ (the factor 2^{j_2} omitted). Note that K_n^{ϖ} is different from $\hat{K}_n^{\sim \varpi}$. The latter is the multiscale window transform of *K*, not a concept of "energy". Another quantity that might be confused with K_n^{ϖ} is $K^{\sim \varpi}$, or the field *K* reconstructed on window ϖ . $K^{\sim \varpi}$ is a property in physical space. It is conceptually different from the phase space-based K_n^{ϖ} for velocity.

Observe that the first term on the right hand side of Eq. (33) is the time change (in 345 difference form) of the kinetic energy on window ϖ at time $2^{-j_2}n$ (scaled by the series 346 length). The second term, which is proportional to $(\Delta t)^2$, is in general very small (of 347 order $O[(\Delta t)^2]$ compared to $\hat{\delta}_n K_n^{\overline{\omega}}$). As shown in Appendix A, it could be significant only 348 when processes with scales of grid size are concerned. Besides, it is expressed in a form 349 of discretized Laplacian. We may thereby view it indistinguishably as a kind of subgrid 350 parameterization and merge it into the dissipation terms. The term $\underline{\hat{\mathbf{v}}}_{n}^{\sim \varpi} \cdot \hat{\delta}_{n} \underline{\hat{\mathbf{v}}}_{n}^{\sim \varpi}$, which is akin to Harrison and Robinson's $\underline{\bar{\mathbf{v}}} \cdot \frac{\partial \bar{\mathbf{v}}}{\partial t}$, is thus merely the change rate of K_{n}^{ϖ} , with a small 351 352 correction of order $(\Delta t)^2$ (t scaled by the series duration). 353

Terms other than $\partial_t \underline{\mathbf{v}}$ and $\partial_t \rho$ in a 3D primitive equation system do not have time derivatives involved. Multiscale window transforms can be applied directly to every field variable in spite of the spatial gradient operators, if any. To continue the derivation, first take a multiscale window transform of (14),

$$\frac{\partial \hat{w}_n^{\sim \varpi}}{\partial z} + \nabla \cdot \underline{\hat{\mathbf{v}}}_n^{\sim \varpi} = 0.$$
(35)

³⁵⁹ Dot product of the momentum equation reconstructed from (13) on window $\overline{\omega}$ with $\hat{\underline{v}}_n^{\infty} \phi_n^{\varrho,j_2}(t)$, followed by an integration with respect to *t* over the domain [0, ϱ), gives the kinetic energy equation for window $\overline{\omega}$. We are now to arrange the right hand side of this equation into a sum of some physically meaningful terms. X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

Look at the pressure work first. By Eq. (35), it is

$$\int_{0}^{\varrho} -\hat{\underline{\mathbf{v}}}_{n}^{\sim\varpi} \cdot \frac{\nabla P^{\sim\varpi}}{\rho_{0}} \phi_{n}^{\varrho,j_{2}}(t) dt$$

$$= -\hat{\underline{\mathbf{v}}}_{n}^{\sim\varpi} \cdot \frac{\nabla \hat{P}_{n}^{\sim\varpi}}{\rho_{0}} = -\frac{1}{\rho_{0}} \left[\nabla \cdot (\hat{P}_{n}^{\sim\varpi} \hat{\underline{\mathbf{v}}}_{n}^{\sim\varpi}) + \frac{\partial}{\partial z} (\hat{P}_{n}^{\sim\varpi} \hat{w}_{n}^{\sim\varpi}) \right] + \hat{w}_{n}^{\sim\varpi} \frac{\partial \hat{P}_{n}^{\sim\varpi}}{\partial z}$$

$$= -\frac{1}{\rho_{0}} \left[\nabla \cdot (\hat{P}_{n}^{\sim\varpi} \hat{\underline{\mathbf{v}}}_{n}^{\sim\varpi}) + \frac{\partial}{\partial z} (\hat{P}_{n}^{\sim\varpi} \hat{w}_{n}^{\infty}) \right] - \frac{g}{\rho_{0}} \hat{w}_{n}^{\sim\varpi} \hat{\rho}_{n}^{\sim\varpi}$$

$$\equiv \Delta_{h} Q_{P_{n}^{\varpi}} + \Delta_{z} Q_{P_{n}^{\varpi}} - b_{n}^{\varpi}, \qquad (36)$$

where $\Delta_h Q_{P_n^{\varpi}}$ and $\Delta_z Q_{P_n^{\varpi}}$ (Q_P the pressure flux) are respectively the horizontal and vertical pressure working rates (Q stands for flux, a convention in many fluid mechanics textbooks). The third term, $-b_n^{\varpi} = -\frac{g}{\rho_0} \hat{w}_n^{\infty \varpi} \hat{\rho}_n^{\infty \varpi}$, is the rate of buoyancy conversion between the kinetic and available potential energies on window ϖ .

Next look at the friction terms $\underline{\mathbf{F}}_{mz}$ and $\underline{\mathbf{F}}_{mh}$ in Eq. (13). They stand for the effect of unresolved sub-grid processes. An explicit expression of them is problem-specific, and is beyond of scope of this paper. We will simply write these two terms as $F_{K^{\varpi},z}$ and $F_{K^{\varpi},h}$, which are related to the $\underline{\mathbf{F}}_{mz}$ and $\underline{\mathbf{F}}_{mh}$ in Eq. (13) as follows:

$$F_{K_n^{\varpi},z} = \underline{\mathbf{\hat{v}}}_n^{\sim \varpi} \cdot (\underline{\widehat{\mathbf{F}}}_{mz})_n^{\sim \varpi}, \qquad (37)$$

 $F_{K_n^{\varpi},h} = \underline{\hat{\mathbf{v}}}_n^{\sim \varpi} \cdot (\underline{\widehat{\mathbf{F}}}_{mh})_n^{\sim \varpi} + (\Delta t)^2 (\hat{\delta}_{n^2}^2 \underline{\hat{\mathbf{v}}}_n^{\sim \varpi} \cdot \hat{\delta}_n \underline{\hat{\mathbf{v}}}_n^{\sim \varpi}).$ (38)

In the above, the correction to $\hat{\delta}_n K_n^{\varpi}$ in (33) has been included, as it behaves like a kind of horizontal dissipation.

For the remaining part, the Coriolis force does not contribute to increase $K_n^{\overline{\omega}}$. The nonlinear terms are what we need to pay attention. Specifically, we need to separate

377
$$\mathrm{NL} = -\widehat{\underline{\mathbf{v}}}_{n}^{\sim \varpi} \cdot \nabla \cdot (\widehat{\underline{\mathbf{v}}} \, \overline{\mathbf{v}})_{n}^{\sim \varpi} - \widehat{\underline{\mathbf{v}}}_{n}^{\sim \varpi} \cdot \frac{\partial}{\partial z} (\widehat{w} \, \overline{\mathbf{v}})_{n}^{\sim \varpi}$$

into two classes of energetics which represent transport and transfer processes, respectively. This can be achieved by performing a decomposition as we did in Section 4 for the 3D case, with the field variable S in (23) replaced by u and v, respectively. Let

$$\underline{\mathbf{Q}}_{h} = \lambda_{c} \underline{\mathbf{\hat{v}}}_{n}^{\sim \varpi} \cdot (\underline{\widehat{\mathbf{v}}} \underline{\mathbf{v}})_{n}^{\sim \varpi} = \lambda_{c} \underline{\mathbf{\hat{v}}}_{n}^{\sim \varpi} \cdot (\underline{\widehat{\mathbf{v}}} \underline{\mathbf{v}})_{n}^{\sim \varpi}, \tag{39}$$

$$Q_z = \lambda_c \underline{\hat{\mathbf{v}}}_n^{\sim \varpi} \cdot (\widehat{w} \underline{\mathbf{v}})_n^{\sim \varpi}, \tag{40}$$

³⁸³ where
$$\lambda_c = \frac{1}{2}$$
. Further define

$$\Delta_h Q_{K_n^{\varpi}} = -\nabla \cdot \underline{\mathbf{Q}}_h, \tag{41}$$

385

2

 $\Delta_z Q_{K_n^{\varpi}} = -\frac{\partial Q_z}{\partial z},\tag{42}$

14

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

$$T_{K_n^{\varpi},h}^* = -\underline{\hat{\mathbf{v}}}_n^{\sim \varpi} \cdot \nabla \cdot (\underline{\widehat{\mathbf{v}}}\underline{\hat{\mathbf{v}}})_n^{\sim \varpi} + \nabla \cdot \underline{\mathbf{Q}}_h, \tag{43}$$

386

387

$$T_{K_n^{\varpi},z}^* = -\hat{\underline{\mathbf{v}}}_n^{\sim \varpi} \cdot \frac{\partial}{\partial z} (\widehat{w} \underline{\widetilde{\mathbf{v}}})_n^{\sim \varpi} + \frac{\partial Q_z}{\partial z}.$$
(44)

388 Then it is easy to show that

389
$$NL = (\Delta_h Q_{K_n^{\varpi}} + \Delta_z Q_{K_n^{\varpi}}) + (T_{K_n^{\varpi},h}^{*} + T_{K_n^{\varpi},z}^{*})$$
(45)

is the transport-transfer separation for which we are seeking, with

$$T_{K_{n}^{\varpi},h}^{*} + T_{K_{n}^{\varpi},z}^{*} = \frac{1}{2} \left[-\hat{\mathbf{v}}_{n}^{\sim \varpi} \cdot \nabla \cdot (\widehat{\mathbf{v}} \, \widehat{\mathbf{v}})_{n}^{\sim \varpi} + \nabla \hat{\mathbf{v}}_{n}^{\sim \varpi} : (\widehat{\mathbf{v}} \, \widehat{\mathbf{v}})_{n}^{\sim \varpi} - \frac{\partial}{\partial z} (\widehat{w} \, \widehat{\mathbf{v}})_{n}^{\sim \varpi} \cdot \hat{\mathbf{v}}_{n}^{\sim \varpi} + \frac{\partial \mathbf{v}}{\partial z} \cdot (\widehat{w} \, \widehat{\mathbf{v}})_{n}^{\sim \varpi} \right]$$
(46)

³⁹⁰ the perfect transfer.

In (45), although $(T^*_{K^{\varpi}_n,h} + T^*_{K^{\varpi}_n,z})$ as a whole is perfect, $T^*_{K^{\varpi}_n,h}$ or $T^*_{K^{\varpi}_n,z}$ alone is not. In order to make them so, introduce the following terms:

$$T_{K_n^{\varpi},h} = T_{K_n^{\varpi},h}^* - \hat{K}_n^{\sim \varpi} \nabla \cdot \hat{\underline{\mathbf{v}}}_n^{\sim \varpi}, \qquad (47)$$

394
$$T_{K_n^{\varpi},z} = T_{K_n^{\varpi},z}^* - \hat{K}_n^{\sim \varpi} \frac{\partial \hat{w}_n^{\sim \varpi}}{\partial z},$$

where \hat{K}_{n}^{∞} is the multiscale window transform of $K = \frac{1}{2} \mathbf{y} \cdot \mathbf{y}$ as a field variable (not K_{n}^{ϖ} , the kinetic energy on window ϖ). Clearly $(T_{K_{n}^{\varpi},h}^{*} + T_{K_{n}^{\varpi},z}^{*}) = (T_{K_{n}^{\varpi},h} + T_{K_{n}^{\varpi},z}^{*})$ by the continuity Eq. (35). It is easy to verify that both $T_{K_{n}^{\varpi},h}$ and $T_{K_{n}^{\varpi},z}$ are perfect transfers using the marginalization property. Decomposition (45) now becomes

NL =
$$(\Delta_h Q_{K_n^{\varpi}} + \Delta_z Q_{K_n^{\varpi}}) + (T_{K_n^{\varpi},h} + T_{K_n^{\varpi},z}).$$
 (49)

In summary, the kinetic energy evolution on window ϖ is governed by

$$\hat{\delta}_{n}K_{n}^{\varpi} = -\nabla \cdot \underline{\mathbf{Q}}_{h} - \frac{\partial Q_{z}}{\partial z} + \left[-\hat{\underline{\mathbf{v}}}_{n}^{\sim \varpi} \cdot \nabla \cdot (\widehat{\underline{\mathbf{v}}}\underline{\mathbf{v}})_{n}^{\sim \varpi} + \nabla \cdot \underline{\mathbf{Q}}_{h} - \hat{K}_{n}^{\sim \varpi} \nabla \cdot \hat{\underline{\mathbf{v}}}_{n}^{\sim \varpi}\right] \\ + \left[-\hat{\underline{\mathbf{v}}}_{n}^{\sim \varpi} \cdot \frac{\partial}{\partial z} (\widehat{w}\underline{\mathbf{v}})_{n}^{\sim \varpi} + \frac{\partial Q_{z}}{\partial z} - \hat{K}_{n}^{\sim \varpi} \frac{\partial \hat{w}_{n}^{\sim \varpi}}{\partial z}\right] - \nabla \cdot \left(\underline{\hat{\mathbf{v}}}_{n}^{\sim \varpi} \frac{\hat{P}_{n}^{\sim \varpi}}{\rho_{0}}\right) \\ - \frac{\partial}{\partial z} \left(\hat{w}_{n}^{\sim \varpi} \frac{\hat{P}_{n}^{\sim \varpi}}{\rho_{0}}\right) - \frac{g}{\rho_{0}} \hat{w}_{n}^{\sim \varpi} \hat{\rho}_{n}^{\sim \varpi} + F_{K_{n}^{\varpi}, z} + F_{K_{n}^{\varpi}, h}, \tag{50}$$

where $\underline{\mathbf{Q}}_h$ and Q_z are defined in (39) and (40). Symbolically this is,

$$\dot{K}_{n}^{\varpi} = \Delta_{h} Q_{K_{n}^{\varpi}} + \Delta_{z} Q_{K_{n}^{\varpi}} + T_{K_{n}^{\varpi},h} + T_{K_{n}^{\varpi},z} + \Delta_{h} Q_{P_{n}^{\varpi}} + \Delta_{z} Q_{P_{n}^{\varpi}} - b_{n}^{\varpi} + F_{K_{n}^{\varpi},z} + F_{K_{n}^{\varpi},h}.$$
(51)

In Appendix D a list of these symbols and their meanings is presented.

DYNAT 708 1-36

(48)

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

6. Multiscale available potential energy equation

To arrive at the multiscale available potential energy equation, take the scale window transform of the time-discretized version of Eq. (16) and multiply it by $c\hat{\rho}_n^{\sim \varpi}$ $(c \equiv g^2/(\rho_0^2 N^2))$. The left hand side becomes, as before,

$$c\hat{\rho}_{n}^{\infty}(\widehat{\delta_{t}\rho})_{n}^{\infty} = c\hat{\rho}_{n}^{\infty}\hat{\delta}_{n}\hat{\rho}_{n}^{\infty} = \hat{\delta}_{n}A_{n}^{\varpi} - (\Delta t)^{2}c(\hat{\delta}_{n}^{2}\hat{\rho}_{n}^{\infty}\hat{\sigma}\cdot\hat{\delta}_{n}\hat{\rho}_{n}^{\infty}),$$

405 where

406
$$A_n^{\varpi} = \frac{1}{2} c (\hat{\rho}_n^{\sim \varpi})^2 = \frac{1}{2} \frac{g^2}{\rho_0^2 N^2} (\hat{\rho}_n^{\sim \varpi})^2$$
(52)

(constant multiplier 2^{j_2} omitted) is the available potential energy at location *n* in the phase space (corresponding to the scaled time $2^{-j_2}n$) for the window ϖ . Compared to $\hat{\delta}_n A_n^{\varpi}$, the correction is of order $(\Delta t)^2$, and could be significant only at small scales, as argued for the kinetic energy case.

For the advection-related terms, the transform followed by a multiplication with $c\hat{\rho}_n^{\sim\varpi}$ yields

$$(\mathrm{AD}) = c\hat{\rho}_n^0 \int_0^{\varrho} \left(-\nabla \cdot (\underline{\mathbf{v}}\rho)^{\sim \varpi} - \frac{\partial (w\rho)^{\sim \varpi}}{\partial z} \right) \phi_n^{\varrho, j_2}(t) \, \mathrm{d}t$$
$$= -c\hat{\rho}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}}\rho)_n^{\sim \varpi} - c\hat{\rho}_n^{\sim \varpi} \frac{\partial}{\partial z} (\widehat{w\rho})_n^{\sim \varpi}.$$

As has been explained in Section 4, we need to collect flux-like terms. In the phase space, these terms are:

$$\Delta_h Q_{A_n^{\varpi}} \equiv -\nabla \cdot [\lambda_c c \hat{\rho}_n^{\sim \varpi} (\widehat{\mathbf{v}} \rho)_n^{\sim \varpi}], \tag{53}$$

414

$$\Delta_z Q_{A_n^{\varpi}} \equiv -\frac{\partial}{\partial z} [\lambda_c c \hat{\rho}_n^{\infty \varpi} (\widehat{w\rho})_n^{\infty \varpi}], \tag{54}$$

where $\lambda_c = \frac{1}{2}$. With this flux representation, (AD) is decomposed as

$$(\mathrm{AD}) = \Delta_h Q_{A_n^{\varpi}} + \Delta_z Q_{A_n^{\varpi}} - [c \hat{\rho}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}} \hat{\rho})_n^{\sim \varpi} + \Delta_h Q_{A_n^{\varpi}}] - \left[c \hat{\rho}_n^{\sim \varpi} \frac{\partial}{\partial z} (\widehat{w} \hat{\rho})_n^{\sim \varpi} + \Delta_z Q_{A_n^{\varpi}} \right].$$

The two brackets as a whole represent a perfect transfer process. However, neither of them alone does so. For physical clarity, we need to make some manipulation.

417 Making use of Eq. (35), and denoting

418
$$TS_{A_n^{\varpi}} \equiv \lambda_c \hat{\rho}_n^{\sim \varpi} (\widehat{w} \rho)_n^{\sim \varpi} \frac{\partial c}{\partial z},$$
(55)

16

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

the above decomposition can be written as

$$(AD) = \Delta_h Q_{A_n^{\varpi}} + \Delta_z Q_{A_n^{\varpi}} - [c\hat{\rho}_n^{\sim\varpi}\nabla\cdot(\widehat{\mathbf{v}}\rho)_n^{\sim\varpi} + \Delta_h Q_{A_n^{\varpi}} - \lambda_c c((\hat{\rho}^2)_n^{\sim\varpi}\nabla\cdot\widehat{\mathbf{v}}_n^{\sim\varpi})] - \left[c\hat{\rho}_n^{\sim\varpi}\frac{\partial}{\partial z}(\widehat{w}\rho)_n^{\sim\varpi} + \Delta_z Q_{A_n^{\varpi}} + TS_{A_n^{\varpi}} - \lambda_c c\left((\hat{\rho}^2)_n^{\sim\varpi}\frac{\partial\widehat{w}_n^{\infty\varpi}}{\partial z}\right)\right] + TS_{A_n^{\varpi}}$$

$$\equiv \Delta_h Q_{A_n^{\varpi}} + \Delta_z Q_{A_n^{\varpi}} + T_{A_n^{\varpi},\partial_h\rho} + T_{A_n^{\varpi},\partial_z\rho} + TS_{A_n^{\varpi}},$$
(56)

where $\Delta_h Q_{A_n^{\overline{w}}}$ and $\Delta_z Q_{A_n^{\overline{w}}}$ are, as we already know, the horizontal and vertical transports. The other pair,

$$T_{A_{n}^{\varpi},\partial_{h}\rho} \equiv -c\hat{\rho}_{n}^{\sim\varpi}\nabla\cdot(\widehat{\mathbf{v}\rho})_{n}^{\sim\varpi} - \Delta_{h}Q_{A_{n}^{\varpi}} + \lambda_{c}c((\widehat{\rho^{2}})_{n}^{\sim\varpi}\nabla\cdot\hat{\mathbf{v}}_{n}^{\sim\varpi})$$
(57)

$$422 T_{A_n^{\varpi},\partial_z\rho} \equiv -c\hat{\rho}_n^{\infty}\frac{\partial}{\partial z}(\widehat{w\rho})_n^{\infty} - \Delta_z Q_{A_n^{\varpi}} - TS_{A_n^{\varpi}} + \lambda_c c\left((\widehat{\rho^2})_n^{\infty}\frac{\partial \widehat{w}_n^{\infty}}{\partial z}\right) (58)$$

represent two perfect transfer processes, as can be easily verified with the definition in
 Section 4.

If necessary, $\Delta_h Q_{A_n^{\varpi}}$ and $T_{A_n^{\varpi},\partial_h\rho}$ can be further decomposed as

$$426 \qquad \Delta_h Q_{A_n^{\varpi}} = \Delta_x Q_{A_n^{\varpi}} + \Delta_y Q_{A_n^{\varpi}}, \tag{59}$$

$$_{427} T_{A_n^{\varpi},\partial_h\rho} = T_{A_n^{\varpi},\partial_x\rho} + T_{A_n^{\varpi},\partial_y\rho},$$

where $\Delta_x Q_{A_n^{\varpi}}$ $(T_{A_n^{\varpi},\partial_x\rho})$ and $\Delta_y Q_{A_n^{\varpi}}$ $(T_{A_n^{\varpi},\partial_y\rho})$ are given by the equation for $\Delta_h Q_{A_n^{\varpi}}$ ($T_{A_n^{\varpi},\partial_h\rho}$) with the gradient operator ∇ replaced by $\partial/\partial x$ and $\partial/\partial y$, respectively.

430 Besides the above fluxes and transfers, there exists an extra term

$$TS_{A_n^{\varpi}} \equiv \lambda_c \hat{\rho}_n^{\sim \varpi} (\widehat{w\rho})_n^{\sim \varpi} \frac{\partial c}{\partial z} = -\lambda_c c \hat{\rho}_n^{\sim \varpi} (\widehat{w\rho})_n^{\sim \varpi} \frac{\partial (\log N^2)}{\partial z}$$
(61)

in the (AD) decomposition (recall $c = g^2/\rho_0^2 N^2$). This term represents an apparent source/sink due to the stationary vertical shear of density, as well as an energy transfer.

Next consider the term $w \frac{N^2 \rho_0}{g}$. Recall that N^2 is a function of z only. It is thus immune to the transform. So

$$c\hat{\rho}_{n}^{\sim \varpi} \frac{\rho_{0}}{g} \cdot (\widehat{wN^{2}})_{n}^{\sim \varpi} = c \frac{N^{2} \rho_{0}}{g} \hat{\rho}_{n}^{\sim \varpi} \hat{w}_{n}^{\sim \varpi} = \frac{g}{\rho_{0}} \hat{w}_{n}^{\sim \varpi} \hat{\rho}_{n}^{\sim \varpi} = b_{n}^{\varpi},$$
(62)

which is exactly the buoyancy conversion between available potential and kinetic energies
on window *σ*.

The diffusion terms are treated the same way as before, they are merely denoted as

441
$$F_{A_n^{\varpi},z} = c\hat{\rho}_n^{\sim \varpi} (\widehat{F_{\rho,z}})_n^{\sim \varpi}, \tag{63}$$

$$F_{A_n^{\varpi},h} = c\hat{\rho}_n^{\infty} \widehat{(F_{\rho,h})}_n^{\infty} + (\Delta t)^2 c(\hat{\delta}_{n^2}^2 \hat{\rho}_n^{\infty} \cdot \hat{\delta}_n \hat{\rho}_n^{\infty}).$$
(64)

Put all the above equations together (with the aid of notations (53), (54) and (61)),

$$\begin{split} \hat{\delta}_n A_n^{\varpi} &= \Delta_h Q_{A_n^{\varpi}} + \Delta_z Q_{A_n^{\varpi}} \\ &+ \left[-c\hat{\rho}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}} \widehat{\rho})_n^{\sim \varpi} - \Delta_h Q_{A_n^{\varpi}} + \lambda_c c((\widehat{\rho^2})_n^{\sim \varpi} \nabla \cdot \hat{\mathbf{v}}_n^{\sim \varpi}) \right] \end{split}$$

(60)

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

$$+ \left[-c\hat{\rho}_{n}^{\infty} \frac{\partial}{\partial z} (\widehat{w\rho})_{n}^{\infty} - \Delta_{z} Q_{A_{n}^{\varpi}} - TS_{A_{n}^{\varpi}} + \lambda_{c} c \left((\widehat{\rho^{2}})_{n}^{\infty} \frac{\partial \widehat{w}_{n}^{\infty}}{\partial z} \right) \right] + TS_{A_{n}^{\varpi}} + \frac{g}{\rho_{0}} \widehat{w}_{n}^{\infty} \widehat{\rho}_{n}^{\infty} + F_{A_{n}^{\varpi},z} + F_{A_{n}^{\varpi},h},$$
(65)

or, in a symbolic form,

DTD 5

$$\dot{A}_{n}^{\varpi} = \Delta_{h} Q_{A_{n}^{\varpi}} + \Delta_{z} Q_{A_{n}^{\varpi}} + T_{A_{n}^{\varpi},\partial_{h}\rho} + T_{A_{n}^{\varpi},\partial_{z}\rho} + TS_{A_{n}^{\varpi}} + b_{n}^{\varpi} + F_{A_{n}^{\varpi},z} + F_{A_{n}^{\varpi},h}.$$
(66)

⁴⁴³ For a list of the meanings of these symbols, refer to Appendix D.

444 7. Horizontal treatment

As in Fourier analysis, the transform coefficients of MWT contain phase information; unlike Fourier analysis, the energies defined in Section 3.2, which are essentially the transform coefficients squared, still contain phase information. This is fundamentally the same as what happens with the real-valued wavelet analysis, which has been well studied in the context of fluid dynamics (e.g., Farge, 1992; Iima and Toh, 1995).

In the presence of advection, the phase information problem leads to superimposed 450 oscillations with high wavenumbers on the spatial distribution of obtained energetics. This 451 may be understood easily, following an argument in the wavelet energetic analysis of shock 452 waves by Iima and Toh (1995). While in the sampling space³ the phase oscillation might not 453 be obvious or even ignored because of the discrete nature in time, in the spatial directions 454 it surfaces through a Galilean transformation. Look at the transform (7). The characteristic 455 frequency is $f_c \sim 2^{j_2}$ cycles over the time duration. (Recall the signals are equally sampled 456 on 2^{j_2} points in time.) Now suppose there is a flow with constant speed u_0 . The oscillation 457 in time with f_c is then transformed to the horizontal plane with a wavelength on the order 458 of u_0/f_c . Suppose the sampling interval is Δt , the time step size for the dataset. Suppose 459 further the spatial grid size is Δx . In a numerical scheme explicit in advection (which is true 460 for most numerical models), it must be smaller than or equal to $\Delta x/u_0$ to satisfy the CFL condition. So the oscillation has a wavenumber $k_c \sim O(\frac{1}{\Delta x})$ or larger, as $f_c \sim \frac{1}{\Delta t}$. Fig. 2a 461 462 shows a typical example of the energetic term for the Iceland-Faeroe Frontal variability (cf. 463 Robinson et al., 1996a,b; LR3). Notice how the substantial energetic information (Fig. 2b) 464 is buried in the oscillations with short wavelengths. (The time sampling interval is $10\Delta t$ 465 here.) 466

The phase oscillation as in Fig. 2a is a technique problem deeply rooted in the nature of localized transforms. It must be eliminated to keep the energetic terms from being blurred. In our case, this is easy to be done. As the characteristic frequency is always 2^{j_2} , the highest for the signal under concern, the oscillation energy peaks at very high wavenumbers, far away from the substantial energy on the spectrum. Except for energetics on the sub-mesoscale

³ Given a scale window, the MWT transform coefficients form a complete function space. We here refer to it as a sampling space.

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx



Fig. 2. (a) The total transfer of APE from the large-scale window to the meso-scale window for the Iceland-Faeroe Frontal variability at depth 300 m on August 21, 1993 (cf. LR3, and Robinson et al., 1996a,b). (b) The horizontally filtered map (units: m^2s^{-3}).

window, a horizontal scaling synthesis with a proper upper scale level (lower enough to 472 avoid the phase problem but higher enough to encompass all the substantial information) 473 will give us all what we want. As a scaling synthesis is in fact a low-pass filtering which 474 may also be loosely understood as a "local averaging", we are taking a measure essentially 475 similar to the time averaging approach of Iima and Toh (1995), except that we are here 476 dealing with the horizontal rather than temporal direction. From now on, all the energetics 477 should be understood to be "locally averaged" with appropriate spatial window bounds, 478 though for notational laconism, we will keep writing them in their original forms. 479

One thing that should be pointed out regarding the MWT is that the phase information 480 to be removed is always located around the highest wavenumbers on the energy spectrum. 481 The reason is that in Eq. (7) a scaling basis at the highest scale level j_2 is used for transforms 482 on all windows. This is in contrast to wavelet analyses, in which the larger the scale for 483 the transform, the larger the scale for the phase oscillation (see Iima and Toh, 1995). The 484 special structure of the MWT transform spectrum is very beneficial to the phase removal. 485 Generally no aliasing will happen in separating the substantial processes from the phase 486 oscillation. 487

8. Connection to the classical formalism

The MS-EVA can be easily connected to a classical energetics formalism, with the aid of the MWT properties presented in Section 2.3, particularly the property of marginalization. For kinetic energy, Appendix C shows that, when

(1) $j_0 = 0$, $j_1 = j_2$ (i.e., only *two-scale windows* are considered), and

493 (2) a periodic extension ($\rho = 1$) is employed,

Eq. (50) for $\varpi = 0$ and $\varpi = 1$ are reduced respectively to the mean and eddy kinetic energy equations in Harrison and Robinson (1978)'s Reynolds-type energetics adapted for open ocean problems [see Eqs. (A.28) and (A.33)]. For available potential energy, the classical

24

ARTICLE IN PRESS

formulation (2D only) in a statistical context gives the following mean and eddy equations
 (e.g., Tennekes and Lumley, 1972)

$$\frac{\partial A_{\text{mean}}}{\partial t} + \nabla \cdot (\underline{\mathbf{v}} A_{\text{mean}}) = -c\bar{\rho}\nabla \cdot \underline{\mathbf{v}'}\rho',\tag{67}$$

$$\frac{\partial A_{\text{eddy}}}{\partial t} + \nabla \cdot \left(\underline{\mathbf{v}}_{2}^{1} c \rho^{2}\right) = -c \overline{\rho' \underline{\mathbf{v}}'} \cdot \nabla \overline{\rho},\tag{68}$$

where $A_{\text{mean}} = \frac{1}{2}c\bar{\rho}^2$, $A_{\text{eddy}} = \frac{1}{2}c\overline{(\rho)'^2}$. Eqs. (67) and (68) can be adapted for open ocean problems by modifying the time rates of change using the approach by Harrison and Robinson (1978). Following the same way as that for KE, these modified equations can be derived directly from the MS-EVA APE Eq. (65) under the above two assumptions.

It is of interest to notice that the multiscale energy Eqs. (50) and (65) appear in the same form for different windows. This is in contrast to the classical Reynolds-type formalism, where the eddy energetics are usually quite different in form from their mean counterparts. This difference disappears if the averaging and deviating operators in (67), (68), (A.28), and (A.33), are rewritten in terms of multiscale window transform. One might have been using the averaging-deviating approach for years without realizing that they actually belong to a kind of transform and synthesis.

Consequently, the classical energetic formalism is equivalent to our MS-EVA under a 512 two-window decomposition with $j_0 = 0$ and $\rho = 1$. The latter can be viewed as a gen-513 eralization of the former for GFD processes occurring on arbitrary scale windows. The 514 MS-EVA capabilities, however, are not limited to this. In (67) and (68), the rhs terms, or 515 transfers as usually interpreted, sum to $-c\nabla \cdot (\bar{\rho}\rho' \mathbf{y}')$, which is generally not zero. That is 516 to say, these "transfers" are not "perfect". They still contain some information of transport 517 processes. Our MS-EVA, in contrast, produces transfers on a different basis. The concept of 518 perfect transfer defined through transfer-transport separation allows us to make physically 519 consistent inference of the energy redistribution through scale windows. In this sense, the 520 MS-EVA has an aspect which is distinctly different from the classical formalism. 521

522 9. Interaction analysis

Different from the classical energetics, a localized energy transfer involves not only 523 interactions between scales, but also interactions between locations in the sampling space. 524 We have already seen this in the definition of perfect transfer processes. A schematic is 525 shown in Fig. 3. The addition of sampling space interaction compounds greatly the transfer 526 problem, as it mingles the inter-scale interactions with transfers within the same scale 527 window, and as a result, useful information tends to be disguised, especially for those 528 processes such as instabilities. We must single out this part in order to have the substantial 529 dynamics up front. 530

In the MS-EVA, transfer terms are expressed in the form of triple products. They are all like

533
$$T(\varpi, n) = \hat{\mathcal{R}}_n^{\sim \varpi}(\widehat{pq})_n^{\sim \varpi}, \quad \text{for } \mathcal{R}, \, p, q \in V_{\varrho, j_2}, \tag{69}$$

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx



Fig. 3. A schematic of the energy transfers toward a meso-scale process at location *n*. Depicted are the transfers from different time scales at the same location (vertical arrows), transfers from surrounding locations at the same scale level (horizontal arrows), and transfers from different scales at different locations (dashed arrows).

a form which we call *basic transfer function* for reference convenience. Using the representation (9), it may be expanded as

536
$$T(\varpi, n) = \sum_{\varpi_1, \varpi_2} \sum_{n_1, n_2} Tr(n, \varpi | n_1, \varpi_1; n_2, \varpi_2),$$
(70)

537 where

53

$$Tr(n, \varpi|n_1, \varpi_1; n_2, \varpi_2) = \hat{\mathcal{R}}_n^{\sim \varpi} \cdot [\hat{p}_{n_1}^{\sim \varpi_1} \hat{q}_{n_2}^{\sim \varpi_2} (\phi_{n_1}^{\widehat{\varrho}, j_2} \phi_{n_2}^{\widehat{\varrho}, j_2})_n^{\sim \varpi}],$$
(71)

and the sums are over all the possible windows and locations. $Tr(n, \varpi|n_1, \varpi_1; n_2, \varpi_2)$ is a *unit expression* of the interaction amongst the triad $(n, w; n_1, w_1; n_2, w_2)$. It stands for the rate of energy transferred to (n, ϖ) from the interaction of (n_1, ϖ_1) and (n_2, ϖ_2) . We will refer to the pairs (n_1, w_1) and (n_2, w_2) as the *giving modes*, and (n, w) the *receiving mode*, a naming convention after Iima and Toh (1995).

Theoretically, expansion of a basic transfer function in terms of unit expression allows one 544 to trace back to all the sources that contributes to the transfer. Practically, however, it is not an 545 efficient way because of the huge number of mode combinations and hence the huge number 546 of triads. In our problem, such a detailed analysis is not at all necessary. If (70) is modified 547 such that some terms are combined, the computational redundancy would be greatly reduced 548 whereas the physical interpretation could be even clearer. We now present the modification. 549 Look at the meso-scale window ($\overline{\omega} = 1$) first. It is of particular importance because it 550 mediates between the large scales and sub-mesoscales on a spectrum. For a field p, make 551 the decomposition 552

$$p = \hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}(t) + p_{*^1} = p^{\sim 0} + \hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}(t) + p_{*^1}^{\sim 1} + p^{\sim 2},$$
(72)

553

 $p_{*1} = p - \hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}(t)$ (73)

DYNAT 708 1-36

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

and $p_{*1}^{\sim 1}$ is the meso-scale part of p_{*1} ,

557

$$p_{*^{1}}^{\sim 1} = p^{\sim 1} - \hat{p}_{n}^{\sim 1} \phi_{n}^{\varrho, j_{2}} = \sum_{i \in \mathcal{N}_{\varrho}^{j_{2}}, i \neq n} \hat{p}_{i}^{\sim 1} \phi_{i}^{\varrho, j_{2}}.$$
(74)

The new interaction analysis concerns the relationship between scales and locations, instead 558 of between triads. The advantage of this is that we do not have to resort to those triad 559 modes, which may not have physical correspondence in the large-scale window, to make 560 interpretation. Note not any $\hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}$ can convincingly characterize $p^{\sim 1}(t)$ at location *n*. But 561 in this context, as the basis function $\phi_n^{\varrho, j_2}(t)$ we choose is a very localized one (localization 562 order delimited, see L02), we expect the removal of $\hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}$ will effectively (though not 563 totally) eliminate from $p^{\sim 1}$ the contribution from location n. This has been evidenced in the 564 example of of a meridional velocity series v (Fig. 4), where at n = 384, $v_{1}^{\sim 1}$ is only about 6% 565 $\left(\left|\frac{-0.0106}{0.17}\right|\right)$ of the $v^{\sim 1}$ in magnitude, while at other locations v and $v_{*^1}^{\sim 1}$ are almost the same 566 (fluctuations negligible around n). Therefore, one may practically, albeit not perfectly, take 567 $\hat{p}_n^{\sim 1} \phi_n^{\varrho, j_2}$ as the meso-scale part of p with contribution from location n only (corresponding to $t = 2^{-j_2}n$), and $p_{*^1}^{\sim 1}$ the part from all locations other than n. Note $p_{*^1}^{\sim 1}$ has an n-dependence. 568 569 For notational clarity, it is suppressed henceforth. 570

Likewise, for field $q \in V_{\varrho, j_2}$, it can also be decomposed as

$$q = q^{-0} + q^{-1} + q^{-2}$$
(75)

572 573

$$q = q^{\sim 0} + \hat{q}_n^{\sim 1} \phi_n^{\varrho, j_2} + q_{*^1}^{\sim 1} + q^{\sim 2},$$
(76)

with interpretation analogous to that of $p_{*1}^{\sim 1}$ for the starred term. The decompositions for p and q yield an analysis of the basic transfer function $T(1, n) = \hat{\mathcal{R}}_n^{\sim 1} \cdot (\widehat{pq})_n^{\sim 1}$ into an interaction matrix, which is shown in Table 1. In this matrix, L stands for large-scale window and S for sub-mesoscale window (all locations). M_n is used to denote the mesoscale contribution from location n, while M_* signifies the meso-scale contributions other than that location. Among these interactions, M_n-M_* and M_*-M_* contribute to T(1, n) from the same scale window (meso-scale, without inter-scale transfers being involved. We may sub-total all the resulting 16 terms into 5 more meaningful terms:

$$T_{n}^{0 \to 1} = \hat{\mathcal{R}}_{n}^{\sim 1} \cdot [(\widehat{p^{\sim 0}q^{\sim 0}})_{n}^{\sim 1} + \hat{q}_{n}^{\sim 1}(\widehat{p^{\sim 0}\varphi_{n}^{\varrho,j_{2}}})_{n}^{\sim 1} + (\widehat{p^{\sim 0}q_{*}^{\sim 1}})_{n}^{\sim 1} + \hat{p}_{n}^{\sim 1}(\widehat{\varphi_{n}^{\varrho,j_{2}}q^{\sim 0}})_{n}^{\sim 1} + (\widehat{p^{\sim 1}q^{\sim 0}})_{n}^{\sim 1}]$$
$$= \hat{\mathcal{R}}_{n}^{\sim 1} \cdot [(\widehat{p^{\sim 0}q^{\sim 0}})_{n}^{\sim 1} + (\widehat{p^{\sim 1}q^{\sim 0}})_{n}^{\sim 1} + (\widehat{p^{\sim 0}q^{\sim 1}})_{n}^{\sim 1}]$$
(77)

$$T_{n}^{2 \to 1} = \hat{\mathcal{R}}_{n}^{\sim 1} \cdot [\hat{p}_{n}^{\sim 1}(\phi_{n}^{\widehat{p},j}q^{\sim 2})_{n}^{\sim 1} + (\widehat{p_{*^{1}}^{(1)}q^{\sim 2}})_{n}^{\sim 1} + \hat{q}_{n}^{\sim 1}(\widehat{p^{\sim 2}\phi_{n}^{\rho},j_{2}})_{n}^{\sim 1} + (\widehat{p^{\sim 2}q_{*^{1}}})_{n}^{\sim 1} + (\widehat{p^{\sim 2}q^{\sim 2}})_{n}^{\sim 1}]$$

$$= \hat{\mathcal{R}}_{n}^{\sim 1} \cdot [(\widehat{p^{\sim 1}q^{\sim 2}})_{n}^{\sim 1} + (\widehat{p^{\sim 2}q^{\sim 2}})_{n}^{\sim 1} + (\widehat{p^{\sim 2}q^{\sim 1}})_{n}^{\sim 1}]$$
(78)

22

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx



Fig. 4. A typical time series of v (in cm/s) from the Iceland-Faeroe Frontal variability simulation (point (35, 43, 2). Refer to Fig. 2 for the location) and its derived series (cf. LR3). There are $2^{j_2} = 1024$ data points, and scale windows are chosen such that $j_0 = 0$ and $j_1 = 4$. The original series v and its large-scale reconstruction $v^{\sim 0}$ are shown in (a), and the meso-scale and sub-mesoscale are plotted in (b) and (c) respectively. Also plotted in (b) is the "starred" series (dotted) $v_{*1}^{\sim 1}$ for location n = 384. (d) is the close-up of (b) around n = 384. Apparently, $v_{*1}^{\sim 1}$ is at least one order smaller than $v^{\sim 1}$ in size at that point, while these two are practically the same at other points. Location n corresponds to a scaled time $t = 2^{-j_2}n$ (here forecast day 8).

574
$$T_n^{0\oplus 2\to 1} = \hat{\mathcal{R}}_n^{\sim 1} \cdot [(\widehat{p^{\sim 2}q^{\sim 0}})_n^{\sim 1} + (\widehat{p^{\sim 0}q^{\sim 2}})_n^{\sim 1}]$$
(79)

$$T_{n \to n}^{1 \to 1} = \hat{\mathcal{R}}_{n}^{\sim 1} \cdot \left[\hat{p}_{n}^{\sim 1} \hat{q}_{n}^{\sim 1} (\widehat{\phi_{n}^{\varrho, j_{2}}})_{n}^{2^{\sim 1}} \right]$$
(80)

$$T_{\text{other}\to n}^{1\to1} = \hat{\mathcal{R}}_n^{\sim 1} \cdot [(\widehat{p^{\sim 1}q_{*^1}})_n^{\sim 1} + \hat{q}_n^{\sim 1}(\widehat{p_{*^1}^{\sim 2}\phi_n^{\varrho,j_2}})_n^{\sim 1}].$$
(81)

Table 1

575

Interaction matrix for basic transfer function $T(1, n)$:	$= \widehat{\mathcal{R}}_n^{\sim 1} \cdot (\widehat{pq})_n^{\sim 1}$
--	--

	$p^{\sim 0}$	$\hat{p}_n^{\sim 1}\phi_n^{arrho,j_2}$	$p_{*^1}^{\sim 1}$	$p^{\sim 2}$
$q^{\sim 0}$	L–L	L-M _n	L-M*	L–S
$\hat{q}_n^{\sim 1} \phi_n^{\varrho, j_2}$	M _n -L	M _n -M _n	M_n-M_*	M _n –S
$q_{\star^1}^{\sim 1}$	M _* -L	M _* -M _n	M_*-M_*	M _* -S
$q^{\hat{\sim}2}$	S–L	S-M _n	S-M*	S–S

ARTICLE IN PRESS

If necessary, $T_{n \to n}^{1 \to 1}$ and $T_{\text{other} \to n}^{1 \to 1}$ may also be combined to one term. The result is denoted as $T_n^{1 \to 1}$.

The physical interpretations of above five terms are embedded in the naming convention of the superscripts, which reveals how energy is transferred to mode (1, n) from other scales. Specifically, $T_n^{0 \rightarrow 1}$ and $T_n^{2 \rightarrow 1}$ are transfer rates from windows 0 and 1, respectively, and $T_n^{0 \oplus 2 \rightarrow 1}$ is the contribution from the window 0–window 2 interaction over the meso-scale range. The last two terms, $T_{n \rightarrow n}^{1 \rightarrow 1}$ and $T_{other \rightarrow n}^{1 \rightarrow 1}$, sum up to $T_n^{1 \rightarrow 1}$, which represents the part of transfer from the same window.

Above are the interaction analysis for T(1, n). Using the same technique, one can obtain a similar analysis for T(0, n) and T(2, n). The results are supplied in Appendix B.

What merits mentioning is that different analyses may be obtained by making different 587 sub-grouping for Eq. (70). The rule of thumb here is to try to avoid those starred terms as 588 in Eq. (81), which makes the major overhead in computation (in terms of either memory or 589 CPU usage). In the above analyses, say the meso-scale analysis, if a whole perfect transfer 590 is calculated, the sum of those terms in the form of $T_{n \to n}^{1 \to 1}$ will vanish by the definition of 591 perfect transfer processes. This also implies that the sum of those transfer functions in the 592 form of $T_{\text{other} \rightarrow n}^{1 \rightarrow 1}$ will be equal to the sum of terms in the same form but with all the stars 593 dropped. Hence in performing interaction analysis for a perfect transfer process, we may 594 simply ignore the stars for the corresponding terms. But if it is an arbitrary transfer term 595 which does not necessarily represent a perfect transfer process (e.g, $TS_{A_n^1}$), the starred-term-596 caused heavy computational overhead will still be a problem. 597

In practice, this overhead may be avoided under certain circumstances. Recall that we 598 have built a highly localized scaling basis function ϕ . For any $p \in V_{\varrho, j_2}$, it yields a function 599 $p(t)\phi_n^{\varrho,j_2}(t)$ with an effective support of the order of the grid size. The large- or meso-600 scale transform of this function is thence negligible, should j_1 be smaller than j_2 by some 601 considerable number (3 is enough). Only when it is in the sub-mesoscale window need 602 we really compute the starred term. An example with a typical time series of ρ and u is 603 plotted in Fig. 5. Apparently, for the large-scale and meso-scale cases, $\hat{\rho}_n^{\sim 0}(u\phi_n^{\varrho,j_2})_n^{\sim 0}$ and 604 $\hat{\rho}_n^{\sim 1}(u\phi_n^{\varrho,\overline{j_2}})_n^{\sim 1}$ (red circles) are very small and hence $(\widehat{\rho^{\sim 0}}_{*^0}u)_n^{\sim 0}$ and $(\widehat{\rho_{*^1}}^{-1}u)_n^{\sim 1}$ can be 605 approximated by $(\rho^{\sim 0}u)_n^{\sim 0}$ and $(\rho^{\sim 1}u)_n^{\sim 1}$, respectively. This approximation fails only in 606 the sub-mesoscale case, where the corresponding two parts are of the same order. 607

It is of interest to give an estimation of the relative importance of all these interaction 608 terms obtained thus far. For the mesoscale transfer function T(1, n), $T_n^{0\oplus 2\to 1}$ is generally 609 not significant (compared to other terms). This is because, on a spectrum, if two processes 610 are far away from each other (as is the case for large scale and sub-mesoscale), they are 611 usually separable and the interaction are accordingly very weak. Even if there exists some 612 interaction, the spawned new processes generally stay in their original windows, seldom 613 going into between. Apart from $T_n^{0\oplus 2\to 1}$, all the others are of comparable sizes, though 614 more often than not $T_n^{0 \to 1}$ dominates the rest (e.g., Fig. 6b). 615

For the large-scale window, things are a little different. This time it is term $T_n^{2\to0}$ that is not significant, with the same reason as above. But term $T_n^{1\oplus2\to0}$ is in general not negligible. In this window, the dominant energy transfer is usually not from other scales, but from other locations at the same scale level. Mathematically this is to say, $T_{other\to n}^{0\to0}$ usually dominates

24 X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx



Fig. 5. An example showing relative importance of the decomposed terms from $T_{A_n^{\infty},\partial_h\rho}$. Data source: same as that in Fig. 4 (zonal velocity only). Units: kg/m²s. Left: $(\widehat{\rho_{*}^{O}u})_n^{\sim 0}$ (heavy solid line) and $\hat{\rho}_n^{\sim 0}(\widehat{u\phi_n^{\rho,j_2}})_n^{\sim 0}$ (circle); middle: $(\widehat{\rho_{*}^{-1}u})_n^{\sim 1}$ (heavy solid line) and $\hat{\rho}_n^{\sim 1}(\widehat{u\phi_n^{\rho,j_2}})_n^{\sim 1}$ (circle); right: $(\widehat{\rho_{*}^{\sim 2}u})_n^{\sim 2}$ (heavy solid line) and $\hat{\rho}_n^{\sim 2}(\widehat{u\phi_n^{\rho,j_2}})_n^{\sim 2}$ (circle). Obviously, the $(\widehat{\rho_{*}^{\circ w}u})_n^{\sim w}$ in the decomposition $(\widehat{\rho^{\sim w}u})_n^{\sim w} = (\widehat{\rho_{*}^{\circ w}u})_n^{\sim w} + \hat{\rho}_n^{\sim w}(\widehat{u\phi_n^{\rho,j_2}})_n^{\sim w}$ can be well approximated by $(\widehat{\rho^{\sim w}u})_n^{\sim w}$ for windows w = 0, 1.

the other terms. This is understandable since a large-scale feature results from interactions
 with modes covering a large range of location on the time series. If each location contributes
 even a little bit, the grand total could be huge. This fact is seen in the example in Fig. 6a.

By the same argument as above, within the sub-mesoscale window, the dominant term is $T_n^{1 \to 2}$. But $T_n^{0 \oplus 1 \to 2}$ could be of some importance also. In comparison to these two, $T_n^{0 \to 2}$ and $T_n^{2 \to 2} = T_{\text{other} \to n}^{2 \to 2} + T_{n \to n}^{2 \to 2}$ are not significant.



Fig. 6. An example showing the relative importance of analytical terms of $T_{K_n^{\varpi},h}$ at 10 (time) locations. The data source and parameter choice are the same as that of Fig. 4. Here the constant factor 2^{j_2} has been multiplied. (a) Analysis of $T_{K_n^0,h}$ (thick solid): $T_{K_n^0,h}^{1\to0}$ (thick dashed), $T_{K_n^0,h}^{2\to0}$ (solid), and $T_{K_n^0,h}^{0\to0}$ (dashed). $T_{K_n^0,h}^{1\oplus2\to0}$ is also shown but unnoticeable. (b) Analysis of $T_{K_n^1,h}$ (thick solid): $T_{K_n^1,h}^{0\to1}$ (thick dashed), $T_{K_n^1,h}^{2\to0}$ (solid), and $T_{K_n^1,h}^{2\to1}$ (solid), and $T_{K_n^1,h}^{1\to1}$ (dashed). $T_{K_n^1,h}^{0\oplus2\to1}$ is also shown but unnoticeable.

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

We finish up this section with two observations of Fig. 6. (1) During the forecast days, $T_{K_n^0,h}$ and $T_{K_n^0,h}^{1\to 0}$ are almost opposite in sign. That is to say, the transfer term without interaction analysis could be misleading in inter-scale energy transfer study. (2) The transfer rates change with time continuously. Analyses in a global time framework apparently do not work here, as application of a global analysis basically eliminates the time structure. This from one aspect demonstrates the advantage of MS-EVA in diagnosing real problems.

632 10. Process classification and energetic scenario

From the above analysis, energetic processes for a geophysical fluid system can be gen-633 erally classified into the following four categories: transport, perfect transfer, buoyancy con-634 version, and dissipation/diffusion. (The apparent source/sink in the multiscale APE equation 635 is usually orders smaller than other terms and hence is negligible.) Dissipation/diffusion is 636 beyond the scope of this paper. All the remaining categories belong to some "conservative" 637 processes. Transport vanishes if integrated over a closed domain; perfect transfer summa-638 rizes to zero over scale windows followed by a marginalization in the sampling space; 639 buoyancy conversion serves as a protocol between the two types of energy. 640

The energetic scenario is now clear. If a system is viewed as defined in a space which 641 includes physical space, phase space, and the space of energy type, then transport, transfer 642 and buoyancy conversion are three mechanisms that redistribute energy through this super 643 space. In a two-window decomposition, communication between the windows are achieved 644 via $T_K^{0\leftrightarrow 1}$ and $T_A^{0\leftrightarrow 1}$. (Here T stands for total transfer, and the superscript $0 \leftrightarrow 1$ for either 645 $0 \rightarrow 1$ or $1 \rightarrow 0$) the two types of energy are converted on each window; while transport 646 brings every point to connection in the physical space. The whole scenario is like an energetic 647 cycle, which is pictorially presented in the left part of Fig. 7 (with all the sub-mesoscale 648 window-related arrows dropped), where arrows are utilized to indicate energy flows, and 649 box and discs for the KE and APE, respectively. 650

When the number of windows increase from 2 to 3, the scenario of energetic processes 651 becomes much more complex. Besides the addition of a sub-mesoscale window, and the 652 corresponding transports, conversions, and the window 1–2 and 0–2 transfers, another pro-653 cess appears. Schematized in Fig. 7 by dashed arrows, it is a transfer to a window from the 654 interaction between another two windows. In traditional jargon, it is a "non-local" transfer, 655 i.e., a transfer between two windows which are not adjacent in the phase space. We do not 656 adopted this language as by "local" in this paper we refer to a physical space context. If the 657 number of windows increases, these "nonlocal" transfers will compound the problem very 658 much, and as a result, the complexity of the energetic scenario will increase exponentially. 659 In a sense, this is one of the reasons why an eddy decomposition is preferred to a wave 660 661 decomposition for multiscale energy study.

662 11. Multiscale enstrophy equation

Vorticity dynamics is an integral part of the MS-EVA. In this section we develop the laws for multiscale enstrophy evolution, which are derived from the vorticity equation. X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx



Fig. 7. A schematic of the multiscale energetics for location *n*. Arrows are used to indicate the energy flow, both in the physical space and phase space, and labeled over these arrows are the processes associated with the flow. The symbols adopted are the same as those listed in Table A.2, except that transport and transfer are the total processes. Interaction analyses are indicated in the superscripts of the *T*-terms, whose interpretation is referred to Section 9. For clarity, transfers from the same window are not shown. From this diagram, we see that transports ($\Delta Q_{R_n^{ar}}, \Delta Q_{P_n^{ar}}, \Delta Q_{A_n^{ar}}$, for windows $\varpi = 0, 1, 2$) occur between different locations in physical space, while transfers (the *T*-terms) mediate between scale windows in phase space. The connection between the two types of energy is established through buoyancy conversion (positive if in the direction as indicated in the parenthesis), which invokes neither scale–scale interactions nor location–location energy exchange.

The equation for vorticity $\zeta = \underline{\mathbf{k}} \cdot \nabla \wedge \underline{\mathbf{v}}$ is obtained by crossing the momentum Eq. (13) followed by a dot product with $\underline{\mathbf{k}}$,

$$\frac{\partial \zeta}{\partial t} = \underline{\mathbf{k}} \cdot \nabla \wedge w \frac{\partial \underline{\mathbf{v}}}{\partial z} - \underline{\mathbf{k}} \cdot \nabla \wedge \left[(f + \zeta) \underline{\mathbf{k}} \wedge \underline{\mathbf{v}} \right] + F_{\zeta, z} + F_{\zeta, h}, \tag{82}$$

where $F_{\zeta,z}$ and $F_{\zeta,h}$ denote respectively the vertical and horizontal diffusion. Making use of the continuity Eq. (14), we get,

$$\frac{\partial \zeta}{\partial t} = \underbrace{-\nabla \cdot (\underline{\mathbf{v}}\zeta) - \frac{\partial}{\partial z}(w\zeta)}_{(I)} \underbrace{-\beta v}_{(II)} \underbrace{+(f+\zeta)\frac{\partial w}{\partial z}}_{(III)} \underbrace{+\underline{\mathbf{k}} \cdot \frac{\partial \underline{\mathbf{v}}}{\partial z} \wedge \nabla w}_{(IV)} \underbrace{+F_{\zeta,z} + F_{\zeta,h}}_{(V)}.$$
(83)

Here $\beta = \partial f / \partial y$ is a constant if a β -plane is approximation is assumed. But in general, it does not need to be so. In Eq. (83), there are five mechanisms that contribute to the change of relative vorticity ζ (e.g., Spall, 1989). Apparently, term (I) is the advection of ζ by the flow, and term (V) the diffusion. β -Effect comes into play through term (II). It is the advection

26

ARTICLE IN PRESS

of planetary vorticity f by meridional velocity v. Vortex tubes may stretch or shrink. The vorticity gain or loss due to stretching or shrinking is represented in term (III). Vortex tube may also tilt. Term (IV) results from such a mechanism.

Enstrophy is the "energy" of vorticity, a positive measure of rotation. It is the square of vorticity: $\mathcal{Z} = \frac{1}{2}\zeta^2$. Following the same practice for multiscale energies, the enstrophy on scale window ϖ at time location *n* is defined as (factor 2^{j_2} omitted for brevity)

$$\mathcal{Z}_n^{\varpi} = \frac{1}{2} (\hat{\zeta}_n^{\sim \varpi})^2.$$
(84)

⁶⁸² The evolution of \mathcal{Z}_n^{ϖ} is derived from Eq. (83).

As before, first discretize the only time derivative term in Eq. (83), $\partial \xi / \partial t$, to $\delta_t \zeta$. Take a multiscale transform of the resulting equation and then multiply it by $\hat{\zeta}_n^{\sim \varpi}$. The left hand side results in the evolution $\hat{\delta}_n Z_n^{\varpi}$ plus a correction term which is of the order Δt^2 , Δt being the time spacing of the series. Merging the correction term into the horizontal diffusion, we get an equation

$$\dot{Z}_{n}^{\varpi} = \underbrace{-\hat{\zeta}_{n}^{\sim\varpi} \left[\nabla \cdot (\underline{\mathbf{v}}\hat{\zeta})_{n}^{\sim\varpi} + \frac{\partial(\widehat{w}\hat{\zeta})_{n}^{\sim\varpi}}{\partial z}\right]}_{(\mathrm{AD})} -\beta\hat{\zeta}_{n}^{\sim\varpi}\hat{v}_{n}^{\sim\varpi} + f\hat{\zeta}_{n}^{\sim\varpi} \left(\frac{\widehat{\partial w}}{\partial z}\right)_{n}^{\sim\varpi} + \hat{\zeta}_{n}^{\sim\varpi}\underline{\mathbf{k}} \cdot \left(\frac{\partial \underline{\mathbf{v}}}{\partial z} \wedge \nabla w\right)_{n}^{\sim\varpi} + F_{Z_{n}^{\varpi},z} + F_{Z_{n}^{\varpi},h}.$$

Again, $F_{Z_n^{\varpi},z}$ and $F_{Z_n^{\varpi},h}$ here are just symbolic representations of the vertical and horizontal diffusions. Following the practice in deriving the APE equation, the process represented by the advection-related terms (AD) can be decomposed into a sum of transport processes and transfer processes. Denote

$$\Delta_h Q_{\mathcal{Z}_n^{\varpi}} = -\nabla \cdot [\lambda_c \hat{\zeta}_n^{\sim \varpi}(\underline{\widehat{\mathbf{v}}}_{\zeta})_n^{\sim \varpi}], \tag{85}$$

688

$$\Delta_z Q_{\mathcal{Z}_n^{\varpi}} = -\frac{\partial}{\partial z} [\lambda_c \hat{\zeta}_n^{\sim \varpi} (\widehat{w} \hat{\zeta})_n^{\sim \varpi}]$$
(86)

then it is

$$\begin{split} \mathbf{A}\mathbf{D} &= \Delta_h Q_{\mathcal{Z}_n^{\varpi}} + \Delta_z Q_{\mathcal{Z}_n^{\varpi}} + \left[-\Delta_h Q_{\mathcal{Z}_n^{\varpi}} - \hat{\boldsymbol{\zeta}}_n^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}} \boldsymbol{\zeta})_n^{\sim \varpi} + \lambda_c (\hat{\boldsymbol{\zeta}}^2)_n^{\sim \varpi} \nabla \cdot \hat{\mathbf{v}}_n^{\sim \varpi} \right] \\ &+ \left[-\Delta_z Q_{\mathcal{Z}_n^{\varpi}} - \hat{\boldsymbol{\zeta}}_n^{\sim \varpi} \frac{\partial (\widehat{w} \boldsymbol{\zeta})_n^{\sim \varpi}}{\partial z} + \lambda_c (\hat{\boldsymbol{\zeta}}^2)_n^{\sim \varpi} \frac{\partial \hat{w}_n^{\sim \varpi}}{\partial z} \right] \\ &\equiv \Delta_h Q_{\mathcal{Z}_n^{\varpi}} + \Delta_z Q_{\mathcal{Z}_n^{\varpi}} + T_{\mathcal{Z}_n^{\varpi}, \partial_h \boldsymbol{\zeta}} + T_{\mathcal{Z}_n^{\varpi}, \partial_z \boldsymbol{\zeta}}, \end{split}$$

where $\Delta_h Q_{\mathcal{Z}_n^{or}}$ and $\Delta_z Q_{\mathcal{Z}_n^{or}}$ represent the horizontal and vertical transports, and $T_{\mathcal{Z}_n^{or},\partial_h\zeta}$, $T_{\mathcal{Z}_n^{or},\partial_z\zeta}$ the transfer rates for two distinct processes. It is easy to prove that both of these processes are perfect transfers. Note the multiscale continuity Eq. (35) has been used in obtaining the above form of decomposition. If necessary, $\Delta_h Q_{\mathcal{Z}_n^{or}}$ and $T_{\mathcal{Z}_n^{or},\partial_h\zeta}$ may be further decomposed into contributions from x and y directions, respectively.

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

The enstrophy equation now becomes, after some algebraic manipulation,

$$\dot{\mathcal{Z}}_{n}^{\varpi} = \Delta_{h} Q_{\mathcal{Z}_{n}^{\varpi}} + \Delta_{z} Q_{\mathcal{Z}_{n}^{\varpi}} + \left[-\Delta_{h} Q_{\mathcal{Z}_{n}^{\varpi}} - \hat{\zeta}_{n}^{\sim \varpi} \nabla \cdot (\widehat{\mathbf{v}} \widehat{\zeta})_{n}^{\sim \varpi} + \lambda_{c} (\zeta^{2})_{n}^{\sim \varpi} \nabla \cdot \widehat{\mathbf{v}}_{n}^{\sim \varpi} \right] \\
+ \left[-\Delta_{z} Q_{\mathcal{Z}_{n}^{\varpi}} - \hat{\zeta}_{n}^{\sim \varpi} \frac{\partial (\widehat{w} \widehat{\zeta})_{n}^{\sim \varpi}}{\partial z} + \lambda_{c} (\widehat{\zeta^{2}})_{n}^{\sim \varpi} \frac{\partial \widehat{w}_{n}^{\sim \varpi}}{\partial z} \right] \\
- \beta \hat{\zeta}_{n}^{\sim \varpi} \hat{v}_{n}^{\sim \varpi} + f \hat{\zeta}_{n}^{\sim \varpi} \frac{\partial \widehat{w}_{n}^{\sim \varpi}}{\partial z} + \hat{\zeta}_{n}^{\sim \varpi} \left(\widehat{\zeta} \frac{\partial \widehat{w}}{\partial z} \right)_{n}^{\sim \varpi} \\
+ \hat{\zeta}_{n}^{\sim \varpi} \underline{\mathbf{k}} \cdot \left(\frac{\partial \widehat{\mathbf{v}}}{\partial z} \wedge \nabla w \right)_{n}^{\sim \varpi} + F_{\mathcal{Z}_{n}^{\varpi}, z} + F_{\mathcal{Z}_{n}^{\varpi}, h}.$$
(87)

Or, symbolically,

$$\dot{\mathcal{Z}}_{n}^{\overline{\omega}} = \Delta_{h} Q_{\mathcal{Z}_{n}^{\overline{\omega}}} + \Delta_{z} Q_{\mathcal{Z}_{n}^{\overline{\omega}}} + T_{\mathcal{Z}_{n}^{\overline{\omega}},\partial_{h}\zeta} + T_{\mathcal{Z}_{n}^{\overline{\omega}},\partial_{z}\zeta} + S_{\mathcal{Z}_{n}^{\overline{\omega}},\beta} + S_{\mathcal{Z}_{n}^{\overline{\omega}},f\nabla\cdot\underline{\mathbf{v}}} + TS_{\mathcal{Z}_{n}^{\overline{\omega}},\zeta\nabla\cdot\underline{\mathbf{v}}} + TS_{\mathcal{Z}_{n}^{\overline{\omega}},\text{tilt}} + F_{\mathcal{Z}_{n}^{\overline{\omega}},z} + F_{\mathcal{Z}_{n}^{\overline{\omega}},h}.$$
(88)

⁶⁹⁴ The meanings of these symbols are tabulated in Appendix D.

Each term of Eq. (88) has a corresponding physical interpretation. We have known 695 that $\Delta_h Q_{\mathcal{Z}_n^{\varpi}}$ and $\Delta_z Q_{\mathcal{Z}_n^{\varpi}}$ are horizontal and vertical transports of \mathcal{Z}_n^{ϖ} , respectively, and 696 $T_{\mathcal{Z}_{n}^{\varpi},\partial_{h}\zeta}$ and $T_{\mathcal{Z}_{n}^{\varpi},\partial_{\tau}\zeta}$ transfer rates for two perfect transfer processes. If ζ is horizontally 697 and vertically a constant, then $T_{\mathbb{Z}_n^{\varpi},\partial_z\zeta}$ and $T_{\mathbb{Z}_n^{\varpi},\partial_h\zeta}$ sum up to zero. We have also explained 698 $F_{\mathbb{Z}_n^{\varpi},z} + F_{\mathbb{Z}_n^{\varpi},h}$ represents the diffusion process. Among the rest terms, $S_{\mathbb{Z}_n^{\varpi},\beta}$ and $S_{\mathbb{Z}_n^{\varpi},f\nabla\cdot\underline{\mathbf{v}}}$ 699 stand for two sources/sinks of \mathcal{Z} due to β -effect and vortex stretching, and $TS_{\mathcal{Z}_n^{\varpi}, \zeta \nabla \cdot \underline{\mathbf{y}}}$ and 700 $TS_{\mathbb{Z}_n^{or}, \text{tilt}}$ transfer as well as generate/destroy enstrophy. Processes cannot be well separated 701 for them. In a 2D system, both $TS_{\mathbb{Z}_n^{\varpi}, \zeta \nabla \cdot \mathbf{v}}$ and $TS_{\mathbb{Z}_n^{\varpi}, \text{tilt}}$ vanish. As a result, the multiscale 702 enstrophy equation is expected to be more useful for a plane flow than for a 3D flow. 703

704 12. Summary and discussion

A new methodology, *multiscale energy and vorticity analysis*, has been developed to investigate the inference of fundamental processes from real oceanic or atmospheric data for complex dynamics which are nonlinear, time and space intermittent, and involve multiscale interactions. Multiscale energy and enstrophy equations have been derived, interpreted, and compared to the energetics in classical formalism.

The MS-EVA is based on a localized orthogonal complementary subspace decomposition. It is formulated with the multiscale window transform, which is constructed to cope with the problem between localization and multiscale representation.⁴ The concept of scale and scale window is introduced, and energy and enstrophy evolutions are then formulated for the large-scale, meso-scale, and sub-mesoscale windows. The formulation is principally in time and hence time scale window, but with a treatment in the horizontal dimension. We emphasize that, before physically interpreted, *all the final energetics should be multiplied by a*

⁴ In the classical framework, multiscale energy does not have location identity of the dimension (time or space) to which the multiscale decomposition is performed.

ARTICLE IN PRESS

constant factor 2^{j_2} , and horizontally filtered with a 2D large-scale window synthesis. When 717 the large-scale window bound $i_0 = 0$, and a periodic extension scheme ($\rho = 1$) is adopted, 718 the multiscale energy Eqs. [(50) and (65)] in a two-window decomposition are reduced to the 719 mean and eddy energy equations in a classical framework. In other words, our MS-EVA is a 720 generalization of the classical energetics formalism to scale windows for generic purposes. 721 We have paid particular attention to the separation of transfers from the energetics re-722 sulting from nonlinearity. The separation is made possible by looking for a special type 723 of process, the so-called perfect transfer. A perfect transfer process carries energy through 724 scale windows, but does not generate nor destroy energy as a whole in the system. 725

Perfect transfer terms can be further decomposed to unravel the complicated windowwindow interactions. This is the so-called interaction analysis. Given a transfer function T, an interaction analysis results in many interaction terms, which can be cast into the following four groups:

 $T^{\pi_1 \to \varpi}, \qquad T^{\varpi_2 \to \varpi}, \qquad T^{\varpi_1 \oplus \varpi_2 \to \varpi}, \qquad T^{\varpi \to \varpi}.$

each characteristic of an interaction process. Here superscripts $\overline{\omega} = 0, 1, 2$ stand for large-, meso-, and sub-meso-scale windows, respectively, and $\overline{\omega}_1 = (\overline{\omega} + 1) \mod 3$, $\overline{\omega}_2 = (\overline{\omega} + 2) \mod 3$. Explicit expressions for these functions are given in Eqs. (77)–(80).

By collecting the MS-EVA terms, energetic processes have been classified into four cate-734 gories: transport, perfect transfer, buoyancy conversion, and dissipation/diffusion processes. 735 Transport vanishes if integrated over a closed physical space; buoyancy conversion medi-736 ates between KE and APE on each individual window; while perfect transfer acts merely to 737 redistribute energy between scale windows. The whole scenario is like a complex cycle, as 738 shown in Fig. 7. These "conservative mechanisms" can essentially make energy reach any-739 where in the super space formed with physical space, phase space, and space of energy type. 740 It is not unreasonable to conjecture that, many patterns generated in geophysical fluid flows, 741 complex as they might appear to be, could be a consequence of these energy redistributions. 742 Our MS-EVA therefore contains energetic information which is fundamental to GFD 743 dynamics. It is expected to provide a useful platform for understanding the complexity of 744 the fluids in which all life on Earth occurs. Direct applications may be set up for investigating 745 the processes of turbulence, wave-current and wave-wave interaction, and the stability for 746 infinite dimensional systems. In the sequels to this paper, we will show how this MS-EVA 747 can be adapted to study a more concrete GFD problem. An avenue to application will be 748 established for localized stability analysis (LR2), and two benchmark stability models will 749 be utilized for validation. In another study (LR3), this methodology will be applied to a real 750 problem to demonstrate how process inference is made easy with otherwise a very intricate 751 dynamical system. 752

753 Acknowledgements

We would like to thank Prof. Donald G.M. Anderson, Dr. Kenneth Brink, and Dr. Arthur
 J. Miller for important and interesting scientific discussions. X. San Liang also thanks

ARTICLE IN PRESS

Dr. Joseph Pedlosky for first raising the issue of transport–transfer separation, and thanks
Prof. Brian Farrell, Prof. Yaneer Bar-Yam, Mr. Wayne Leslie, Dr. Patrick Haley, Dr. Pierre
Lermusiaux, Dr. Carlos Lozano, Ms. Gioia Sweetland and Dr. James Wang for their generous
help. This work was supported by the Office of Naval Research under Contracts N0001495-1-0371, N00014-02-1-0989 and N00014-97-1-0239 to Harvard University.

761 Appendix A. Correction to the time derivative term

We have shown in Section 5 that there exists a correction term in the formulas with time
 derivatives. For a kinetic equation, this formula is

$$\underbrace{\hat{\delta}_n K_n}_{(K)} - \underbrace{(\Delta t)^2 (\hat{\delta}_{n2}^2 \hat{\underline{\mathbf{v}}}_n \cdot \hat{\delta}_n \hat{\underline{\mathbf{v}}}_n)}_{(C)}, \tag{A.1}$$

where (C) is the correction term. Scale superscripts are omitted here since we do not want to limit the discussion to any particular scale window. Let's first do some nondimensional analysis so that a comparison is possible. Scale $\hat{\underline{v}}_n$ with U, t with T, then

Term (K)
$$\sim \frac{U^2}{T}$$
, Term (C) $\sim (\Delta t)^2 \frac{U}{T^2} \cdot \frac{U}{T} = (\Delta t)^2 \frac{U^2}{T^3}$.

This enables us to evaluate the weight of (C) relative to (K):

770
$$\frac{\operatorname{Term}\left(\mathrm{C}\right)}{\operatorname{Term}\left(\mathrm{K}\right)} \sim \frac{(\Delta t)^2 U^2 / T^3}{U^2 / T} = \left(\frac{\Delta t}{T}\right)^2.$$

Apparently, this ratio will become significant only when $T \sim \Delta t$, i.e., when the time scale r72 is of the time step size. In our MS-EVA formulation, the correction term (C) is hence not



Fig. A.1. $\hat{\delta}_n K_n$ (thick solid) and its correction term (dashed) for the large-scale (left), meso-scale (middle), and sub-mesoscale (right) kinetic energy equations. Data source and parameter choice are the same as those of Fig. 4 (units in m²/s³; factor 2^{j₂} not multiplied).

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

significant for both large-scale and meso-scale equations. Fig. A.1 confirms this conclusion.
The correction (dashed line) is so small in either the left or middle plots that it is totally
negligible. Only in the sub-mesoscale window can its effect be seen, which, as argued
before, might be parameterized into the dissipation/diffusion.

Appendix B. Interaction analysis for T(0, n) and T(2, n)

Using the technique same as that for T(1, n) in Section 9, we obtain a similar analysis for T(0, n):

$$T(0,n) = \hat{\mathcal{R}}_n^{\sim 0} \cdot (\widehat{pq})_n^{\sim 0} = T_n^{1 \to 0} + T_n^{2 \to 0} + T_n^{1 \oplus 2 \to 0} + T_{n \to n}^{0 \to 0} + T_{\text{other} \to n}^{0 \to 0}, \quad (A.2)$$

782
$$T_n^{1\to0} = \hat{\mathcal{R}}_n^{\sim 0} \cdot [(\widehat{p^{\sim 1}q^{\sim 1}})_n^{\sim 0} + (\widehat{p^{\sim 1}q^{\sim 0}})_n^{\sim 0} + (\widehat{p^{\sim 0}q^{\sim 1}})_n^{\sim 0}]$$
(A.3)

783
$$T_n^{2\to0} = \hat{\mathcal{R}}_n^{\sim 0} \cdot [(\widehat{p^{\sim 0}q^{\sim 2}})_n^{\sim 0} + (\widehat{p^{\sim 2}q^{\sim 2}})_n^{\sim 0} + (\widehat{p^{\sim 2}q^{\sim 0}})_n^{\sim 0}]$$
(A.4)

784
$$T_n^{1\oplus 2\to 0} = \hat{\mathcal{R}}_n^{\sim 0} \cdot [(p^{\sim 2}q^{\sim 1})_n^{\sim 0} + (p^{\sim 1}q^{\sim 2})_n^{\sim 0}]$$
(A.5)

785
$$T_{n \to n}^{0 \to 0} = \hat{\mathcal{R}}_{n}^{\sim 0} \cdot [\hat{p}_{n}^{\sim 0} \hat{q}_{n}^{\sim 0} (\widehat{\phi_{n}^{\varrho, j_{2}}})_{n}^{2^{\circ 0}}]$$
(A.6)

786
$$T_{\text{other}\to n}^{0\to0} = \hat{\mathcal{R}}_n^{\sim 0} \cdot [(p^{\sim 0}q^{\sim 0}_{*^0})_n^{\sim 0} + \hat{q}_n^{\sim 0}(p^{\sim 0}_{*^0}\phi_n^{\varrho,j_2})_n^{\sim 0}], \tag{A.7}$$

787 and T(2, n):

$$T(2,n) = \hat{\mathcal{R}}_n^{\sim 2} \cdot (\widehat{pq})_n^{\sim 2} = T_n^{0 \to 2} + T_n^{1 \to 2} + T_n^{0 \oplus 1 \to 2} + T_{n \to n}^{2 \to 2} + T_{other \to n}^{2 \to 2}, \quad (A.8)$$

789 where

90
$$T_n^{0 \to 2} = \hat{\mathcal{R}}_n^{\sim 2} \cdot [(\widehat{p^{\sim 0}q^{\sim 0}})_n^{\sim 2} + (\widehat{p^{\sim 2}q^{\sim 0}})_n^{\sim 2} + (\widehat{p^{\sim 0}q^{\sim 2}})_n^{\sim 2}]$$
(A.9)

791
$$T_n^{1 \to 2} = \hat{\mathcal{R}}_n^{\sim 2} \cdot [(\widehat{p^{\sim 1} q^{\sim 2}})_n^{\sim 2} + (\widehat{p^{\sim 1} q^{\sim 1}})_n^{\sim 2} + (\widehat{p^{\sim 2} q^{\sim 1}})_n^{\sim 2}]$$
(A.10)

7

$$T_n^{0\oplus 1\to 2} = \hat{\mathcal{R}}_n^{\sim 2} \cdot [(\widehat{p^{\sim 0}q^{\sim 1}})_n^{\sim 2} + (\widehat{p^{\sim 1}q^{\sim 0}})_n^{\sim 2}]$$
(A.11)

$$T_{n \to n}^{2 \to 2} = \hat{\mathcal{R}}_{n}^{\sim 2} \cdot [\hat{p}_{n}^{\sim 2} \hat{q}_{n}^{\sim 2} (\widehat{\phi_{n}^{\varrho, j_{2}}})_{n}^{2^{\sim 2}}]$$
(A.12)

794
$$T_{\text{other}\to n}^{2\to2} = \hat{\mathcal{R}}_n^{\sim 2} \cdot [(\widehat{p^{\sim 2}q_{*^2}^{\sim 2}})_n^{\sim 2} + \hat{q}_n^{\sim 2}(\widehat{p_{*^2}^{\sim 2}\phi_n^{\varrho,j_2}})_n^{\sim 2}].$$
(A.13)

⁷⁹⁵ In these analyses,
$$p_{*^0}$$
 and p_{*^2} are defined as

796
$$p_{*^0} = p - \hat{p}_n^{\sim 0} \phi_n^{\varrho, j_2}(t),$$
 (A.14)

797
$$p_{*2} = p - \hat{p}_n^{\sim 2} \phi_n^{\varrho, j_2}(t).$$
 (A.15)

The physical meaning of the interaction terms is embedded in these mnemonic notations.
 In the superscripts, arrows signify the directions of energy transfer and the numbers 0–2 represent the large-scale, meso-scale, and sub-mesoscale windows, respectively.

ARTICLE IN PRESS

32 X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

Appendix C. Connection between the MS-EVA KE equations and the mean and eddy KE equations in a classical Reynolds formalism

To connect our MS-EVA to the classical energetic formulation, rewrite Eq. (50) (dissipation omitted) as

$$\hat{\mathbf{y}}_{n}^{\sim\varpi} \cdot \left(\frac{\widehat{\partial}\underline{\mathbf{y}}}{\partial t}\right)_{n}^{\sim\varpi} = \underbrace{\hat{\mathbf{y}}_{n}^{\sim\varpi} \cdot \left[-\nabla \cdot (\widehat{\mathbf{y}}\,\widehat{\mathbf{y}})_{n}^{\sim\varpi} - \frac{\partial}{\partial z} (\widehat{w}\,\widehat{\mathbf{y}})_{n}^{\sim\varpi}\right]}_{(\mathbf{I})} + \underbrace{\Delta_{h}Q_{P_{n}^{\varpi}} + \Delta_{z}Q_{P_{n}^{\varpi}} - b_{n}^{\varpi}}_{(\mathbf{I})},$$
(A.16)

We want to see what this equation reduces to if $j_1 = j_2$ (that is to say, only *two-scale windows* are considered), $j_0 = 0$, and a *periodic extension* is employed.

First consider the large scale window $\varpi = 0$. Let *q* be any field variable (*u*, *v*, *w*, or *P*). A two-scale window decomposition means

806
$$q = q^{\sim 0} + q^{\sim 1}$$
. (A.17)

With the choice of zero j_0 and periodic extension, we know from the MWT properties (see Section 2.3) that $q^{\sim 0}$ is constant in time and is equal to \bar{q} or $2^{j_2/2} \hat{q}_n^{\sim 0}$ in magnitude, that is,

⁸¹⁰
$$q^{\sim 0} = \bar{q} = 2^{j_2/2} \hat{q}_n^{\sim 0}, \qquad q^{\sim 1} = q - \bar{q} = q'.$$
 (A.18)

811 Hence

812

$$(\hat{q})_n^{\sim 0} = (\widehat{q^{\sim 0}})_n^{\sim 0} = q^{\sim 0} = 2^{-j_2/2}\bar{q},$$
(A.19)

813
$$(\widehat{q}')_n^{\sim 0} = (\widehat{q^{\sim 1}})_n^{\sim 0} = 0.$$
 (A.20)

Substituting $\underline{\mathbf{v}}$ and w for the q in (A.17), the velocity field is decomposed as $\underline{\mathbf{v}} = \underline{\mathbf{v}} + \underline{\mathbf{v}}'$, and $w = \overline{w} + w'$. Let $K_{\text{mean}} = \frac{1}{2} \overline{\mathbf{v}} \cdot \overline{\mathbf{v}}$. The equivalence between the large-scale transform and duration average allows an expression of the large-scale kinetic energy K_n^0 in terms of K_{mean} . In fact,

⁸¹⁸
$$K_n^0 = 2^{j_2} \left(\frac{1}{2} \underline{\hat{\mathbf{v}}}_n^{\sim 0} \cdot \underline{\hat{\mathbf{v}}}_n^{\sim 0} \right) = \frac{1}{2} \underline{\bar{\mathbf{v}}} \cdot \underline{\bar{\mathbf{v}}} = K_{\text{mean}}.$$
(A.21)

Note here we have taken into account the multiplier 2^{j_2} . These facts are now used to simplify the term (**I**) of Eq. (A.16). With the two-scale decomposition, the dyad ($\underline{\mathbf{v}} \, \underline{\mathbf{v}}$) after transform is expanded as

$$(\widehat{\mathbf{v}}\,\widehat{\mathbf{v}})_{n}^{\sim 0} = (\widehat{\overline{\mathbf{v}}}\,\widehat{\overline{\mathbf{v}}})_{n}^{\sim 0} + (\widehat{\overline{\mathbf{v}}}\,\widehat{\mathbf{v}'})_{n}^{\sim 0} + (\widehat{\mathbf{v}'}\,\widehat{\overline{\mathbf{v}}})_{n}^{\sim 0} + (\widehat{\mathbf{v}'}\,\underline{\mathbf{v}'})_{n}^{\sim 0}$$
(A.22)

$$(\widehat{\mathbf{v}}\widehat{\mathbf{v}})_{n}^{\sim 0} = \underline{\tilde{\mathbf{v}}}_{n}^{\sim 0} + (\widehat{\mathbf{v}'}\underline{\mathbf{v}'})_{n}^{\sim 0}.$$
(A.23)

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

824 Likewise,

825

$$(\widehat{w}\underline{\mathbf{v}})_n^{\sim 0} = \bar{w}\underline{\mathbf{\hat{v}}}_n^{\sim 0} + (\widehat{w'\underline{\mathbf{v}}'})_n^{\sim 0}.$$
(A.24)

These allow term (\mathbf{I}) to be written as

$$(\mathbf{I}) = \hat{\underline{\mathbf{v}}}_{n}^{\sim 0} \cdot \left[-\nabla \cdot (\bar{\underline{\mathbf{v}}} \, \hat{\underline{\mathbf{v}}}_{n}^{\sim 0}) - \frac{\partial}{\partial z} (\bar{w} \, \hat{\underline{\mathbf{v}}}_{n}^{\sim 0}) \right] + \hat{\underline{\mathbf{v}}}_{n}^{\sim 0} \cdot \left[-\nabla \cdot (\bar{\underline{\mathbf{v}}} \, \underline{\mathbf{v}})_{n}^{\sim 0} - \frac{\partial}{\partial z} (\bar{w} \, \underline{\mathbf{v}})_{n}^{\sim 0} \right]$$
$$= 2^{-j_{2}} \left\{ -\nabla \cdot (\bar{\underline{\mathbf{v}}} K_{\text{mean}}) - \frac{\partial}{\partial z} (\bar{w} K_{\text{mean}}) + \bar{\underline{\mathbf{v}}} \cdot \left[-\nabla \cdot (\overline{\underline{\mathbf{v}}' \, \underline{\mathbf{v}}'}) - \frac{\partial}{\partial z} (\overline{w' \, \underline{\mathbf{v}}'}) \right] \right\}$$
$$= 2^{-j_{2}} \left\{ -\nabla \cdot (\bar{\underline{\mathbf{v}}} K_{\text{mean}}) - \frac{\partial}{\partial z} (\bar{w} K_{\text{mean}}) + \bar{\underline{\mathbf{v}}} \cdot \nabla_{3} \cdot \underline{\underline{\mathbf{T}}} \right\}, \tag{A.25}$$

826 where

 $\nabla_3 = \underline{\mathbf{i}}\frac{\partial}{\partial x} + \underline{\mathbf{j}}\frac{\partial}{\partial y} + \underline{\mathbf{k}}\frac{\partial}{\partial z},$

828 and

82

$$= \underbrace{\mathbf{\underline{T}}}_{=} \begin{bmatrix} -\overline{(u'u')} & -\overline{(u'v')} & -\overline{(u'w')} \\ -\overline{(v'u')} & -\overline{(v'v')} & -\overline{(v'w')} \\ -\overline{(w'u')} & -\overline{(w'v')} & -\overline{(w'w')} \end{bmatrix}.$$
(A.26)

⁸³⁰ For term (II), it is equal to, in the present setting,

$$(\mathbf{II}) = 2^{-j_0} \left\{ -\frac{1}{\rho_0} \nabla \cdot (\bar{P} \underline{\bar{v}}) - \frac{1}{\rho_0} \frac{\partial}{\partial z} (\bar{P} \bar{w}) - \frac{g}{\rho_0} \bar{w} \bar{\rho}. \right\}$$
(A.27)

Substitute (I) and (II) back to Eq. (A.16). Considering that the left hand side is now $2^{-j_0} \overline{\underline{v}} \cdot \overline{\left(\frac{\partial \underline{v}}{\partial t}\right)}$, we have, with the common factor 2^{-j_0} cancelled out,

$$\underline{\bar{\mathbf{v}}} \cdot \overline{\left(\frac{\partial \underline{\mathbf{v}}}{\partial t}\right)} = -\nabla \cdot (\underline{\bar{\mathbf{v}}}K^L) - \frac{\partial}{\partial z} (\bar{w}K^L) - \frac{1}{\rho_0} \nabla \cdot (\bar{P}\underline{\bar{\mathbf{v}}}) - \frac{1}{\rho_0} \frac{\partial}{\partial z} (\bar{P}\bar{w}) - \frac{g}{\rho_0} \bar{w}\bar{\rho}. + \underline{\bar{\mathbf{v}}} \cdot \nabla_3 \cdot \underline{\underline{\mathbf{T}}}.$$
(A.28)

This is exactly what Harrison and Robinson (1978) have obtained for the mean kinetic energy, with $\underline{\mathbf{T}}$ the Reynolds stress tensor in their formulation.

Above is about the large-scale energetics. For the meso-scale window ($\varpi = 1$), things are more complicated. In order to make Eq. (A.16) comparable to the classical eddy KE equation, just $j_0 = 0$ and periodic extension are not enough, as now there no longer exists for variable p a linear relation between $\hat{p}_n^{\sim 1}$ and p'. We have to marginalize (A.16) to the physical space to fulfill this mission. In this particular case, the marginalization equality

ARTICLE IN PRESS

34 X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

838 (11) in Section 2.3 is simply

$$\mathcal{M}_n \hat{p}_n^{\sim 1} \hat{q}_n^{\sim 1} = \overline{p'q'}, \tag{A.29}$$

since here the deviation operation (prime) and the meso-scale synthesis operator are identical. Marginalization of (A.16) with $\varpi = 1$ yields

$$\mathbf{\underline{v}'} \cdot \left(\frac{\partial \mathbf{\underline{v}}}{\partial t}\right)' = -\underbrace{\mathbf{\underline{v}'} \cdot \nabla \cdot (\mathbf{\underline{vv}})'}_{(\mathbf{I}')} - \underbrace{\mathbf{\underline{v}'} \cdot \frac{\partial}{\partial z} (w\mathbf{\underline{v}})'}_{(\mathbf{II}')} - \underbrace{\mathbf{\underline{v}'} \cdot \nabla \left(\frac{P'}{\rho_0}\right)}_{(\mathbf{III}')}.$$
(A.30)

⁸⁴³ It is easy to show, as we did before,

(III') =
$$\nabla \cdot \overline{\left(\underline{\mathbf{v}}' \frac{P'}{\rho_0}\right)} + \frac{\partial}{\partial z} \overline{\left(w' \frac{P'}{\rho_0}\right)} + \frac{g}{\rho_0} \overline{w' \rho'}.$$
 (A.31)

⁸⁴⁵ The other two terms sum up to

⁸⁴⁶
$$(\mathbf{I}') + (\mathbf{II}') = \nabla \cdot \overline{\left(\underline{\mathbf{v}} \, \underline{\mathbf{v}'} \cdot \underline{\mathbf{v}'}\right)} + \frac{\partial}{\partial z} \overline{\left(w \, \underline{\mathbf{v}'} \cdot \underline{\mathbf{v}'}\right)} + \overline{\underline{\mathbf{v}'} \underline{\mathbf{v}'}} : \nabla \overline{\underline{\mathbf{v}}} + \overline{\underline{\mathbf{v}'} w'} \cdot \frac{\partial \overline{\underline{\mathbf{v}}}}{\partial z}.$$
 (A.32)

Therefore,

$$\overline{\underline{\mathbf{v}}' \cdot \left(\frac{\partial \underline{\mathbf{v}}}{\partial t}\right)'} = -\nabla \cdot \overline{\left(\underline{\mathbf{v}}\frac{\underline{\mathbf{v}}' \cdot \underline{\mathbf{v}}'}{2}\right)} - \frac{\partial}{\partial z} \overline{\left(w\frac{\underline{\mathbf{v}}' \cdot \underline{\mathbf{v}}'}{2}\right)} - \nabla \cdot \overline{\left(\underline{\mathbf{v}}'\frac{P'}{\rho_0}\right)} - \frac{\partial}{\partial z} \overline{\left(w'\frac{P'}{\rho_0}\right)} - \frac{g}{\rho_0} \overline{w'\rho'} - \overline{\underline{\mathbf{v}}'\underline{\mathbf{v}}'} : \nabla \overline{\underline{\mathbf{v}}} - \overline{\underline{\mathbf{v}}'w'} \cdot \frac{\partial \overline{\underline{\mathbf{v}}}}{\partial z}.$$
(A.33)

Again, this is exactly the eddy KE equation obtained by Harrison and Robinson (1978).

848 Appendix D. Glossary

849 Tables A.1–A.3.

Table A.1

General symbols		
A_n^{ϖ}	Available potential energy on window ϖ at time $2^{-j_2}n$	
j_0, j_1, j_2	Upper bounds of scale level for the three scale windows	
$K_n^{\overline{\omega}}$	Kinetic energy on window ϖ at time $2^{-j_2}n$	
V_{ϱ, j_2}	Direct sum of the three scale windows.	
$\overline{\omega}$	Window index ($\varpi = 0, 1, 2$ for large-scale, meso-scale, and sub-mesoscale windows, respectively)	
\mathcal{Z}_n^{ϖ}	Enstrophy on window ϖ at time $2^{-j_2}n$	
$\hat{z}_n^{\sim \varpi}$	Multiscale window transform of variable z	
$z^{\sim \overline{\omega}}$	Multiscale window synthesis of variable z	
z	Duration average of variable z	
$\phi_n^{\varrho,j}$	Periodized scaling basis function at level j	
$\psi_n^{\varrho,j}$	Periodized wavelet basis function at level j	

ARTICLE IN PRESS

X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx

Kinetic energy (KE)		Available potential energy (APE)	
$\dot{K}_n^{\overline{w}}$	Time rate of change of KE	\dot{A}_n^{ϖ}	Time rate of change of APE
$\Delta_h Q_{K_n^{\overline{w}}}$	Horizontal KE advective working rate	$\Delta_h Q_{A_n^{\overline{w}}}$	Horizontal APE advective working rate
$\Delta_z Q_{K_n^{\varpi}}$	Vertical KE advective working rate	$\Delta_z Q_{A_n^{\varpi}}$	Vertical APE advective working rate
$T_{K_n^{\overline{w}},h}$	Rate of KE transfer due to the hor- izontal flow	$T_{A_n^{\overline{w}},\partial_h \rho}$	Rate of APE transfer due to the horizontal gradient density
$T_{K_n^{\overline{w}},z}$	Rate of KE transfer due to the ver- tical flow	$T_{A_n^{\overline{w}},\partial_z \rho}$	Rate of APE transfer due to the vertical gradient density
$-b_n^{\varpi}$	Rate of buoyancy conversion	b_n^{ϖ}	Rate of inverse buoyancy conversion
$\Delta_h Q_{P_n^{\overline{\omega}}}$	Horizontal pressure working rate	$TS_{A_n^{\varpi}}$	Rate of an imperfect APE transfer due to the stationary shear of the density profile
$\Delta_z Q_{P_n^{\overline{o}}}$	Vertical pressure working rate	$F_{A_n^{\overline{\omega}},h}$	Horizontal diffusion
$F_{K_n^{\overline{w}},z}$ $F_{K_n^{\overline{w}},h}$	Vertical dissipation Horizontal dissipation	$F_{A_n^{\overline{w}},z}$	Vertical diffusion

Table A.2		
Symbols for the multiscale energy equation	ns (time $2^{-j_2}n$, window ϖ)

Table A.3

Symbols for the multiscale enstrophy equation (time $2^{-j_2}n$, window $\overline{\omega}$)

$\dot{Z}_n^{\overline{\omega}}$	Time rate of change of \mathcal{Z} on window ϖ at time $2^{-j_2}n$
$\Delta_h Q_{\mathbb{Z}_n^{\overline{\omega}}}$	Horizontal transport rate
$\Delta_z Q_{\mathcal{Z}_n^{\overline{w}}}$	Vertical transport rate
$T_{\mathcal{Z}_{n}^{\varpi},\partial_{h}\zeta}$	Rate of enstrophy transfer due to the horizontal variation of ζ
$T_{\mathcal{Z}_n^{\overline{\alpha}},\partial_z\zeta}$	Rate of enstrophy transfer due to the vertical variation of ζ
$S_{\mathcal{Z}_n^m,\beta}$	β -Effect-caused source/sink
$S_{\mathcal{Z}_n^{\overline{\alpha}}, f \nabla \cdot \mathbf{v}}$	Source/sink of enstrophy due to horizontal divergence
$TS_{\mathbb{Z}_n^{\overline{m}}, \zeta \nabla \cdot \mathbf{v}}$	Rate of \mathcal{Z} transfer and generation due to rotation-divergence correlation
$TS_{Z_{\alpha}^{\overline{\alpha}}, \text{tilt}}$	Rate of \mathcal{Z} transfer and generation due to the vortex tube tilting
$F_{\mathcal{Z}_n^m,h}$	Horizontal diffusion rate
$F_{Z_n^{m},z}$	Vertical diffusion rate

850 **References**

- Cronin, M., Watts, R., 1996. Eddy-mean flow interaction in the Gulf Stream at 68°W. Part I: Eddy energetics. J.
 Phys. Oceanogr. 26, 2107–2131.
- Farge, M., 1992. Wavelet transforms and their applications to turbulence. Annu. Rev. Fluid Mech. 24, 395–457.
- Fournier, A., 1999. Atmospheric energetics in the wavelet domain I: Governing equations and interpretation for
 idealized flows. Ph.D. Thesis, University of Maryland.
- Harrison, D.E., Robinson, A.R., 1978. Energy analysis of open regions of turbulent flows-mean eddy energetics
 of a numerical ocean circulation experiment. Dyn. Atmos. Oceans 2, 185–211.
- Hernández, E., Weiss, G., 1996. A First Course on Wavelets. CRC Press, 489 pp.
- Holland, W.R., Lin, L.B., 1975. On the generation of mesoscale eddies and their contribution to the ocean general
 circulation. I. A preliminary numerical experiment. J. Phys. Oceangr. 5, 642–657.
- Huang, N.E., Shen, Z., Long, S.R., 1999. A new view of nonlinear water waves: The Hilbert spectrum. Annu. Rev.
 Fluid Mech. 31, 417–457.
- Iima, M., Toh, S., 1995. Wavelet analysis of the energy transfer caused by convective terms: Application to the
 Burgers shock. Phys. Rev. E 52 (6), 6189–6201.

ARTICLE IN PRESS

- 36 X. San Liang, A.R. Robinson / Dynamics of Atmospheres and Oceans xxx (2005) xxx-xxx
- Liang, X.S., 2002. Wavelet-based multiscale window transform and energy and vorticity analysis. Ph.D. Thesis,
 Harvard University, Cambridge, MA, 411 pp.
- Liang, X.S., Anderson, D.G.M., 2003. Multiscale window analysis, in preparation.
- Liang, X.S., Robinson, A.R., 2003a. Multiscale energy and vorticity analysis. II. Instability theory and validation,
 Dyn. Atmos. Oceans (this issue).
- Liang, X.S., Robinson, A.R., 2003b. A study of the Iceland-Faeroe Frontal variability with the multiscale energy
 and vorticity analysis, J. Phys. Oceanogr.
- Liang, X.S., Robinson, A.R., 2003c. Localized stability analysis and its application, J. Fluid Mech., in preparation.
- Pinardi, N., Robinson, A.R., 1986. Quasigeostrophic energetics of open ocean regions. Dyn. Atmos. Oceans 10,
- 874 185–219.
- Plumb, R.A., 1983. A new look at the energy cycle. J. Atmos. Sci. 40, 1669–1688.
- Robinson, A.R., Arango, H.G., Miller, A.J., Warn-Varnas, A., Poulain, P.-M., Leslie, W.G., 1996a. Real-time
 operational forecasting on shipboard of the Iceland-Faeroe Frontal variability. Bull. Am. Meteorol. Soc., 243–
 259.
- 879 Robinson, A.R., Arango, H.G., Warn-Varnas, A., Leslie, W., Miller, A.J., Haley, P., Lozano, C., 1996b. Real-
- time regional forecasting. In: Malanotte-Rizzoli, P. (Ed.), Modern Approaches to Data Assimilation in Ocean
 Modeling. Elsevier Oceanogr. Ser., vol. 61., pp. 377–410.
- Spall, M., 1989. Regional primitive equation modeling and analysis of the POLYMODE data set. Dyn. Atmos.
 Oceans 14, 125–174.
- Tennekes, H., Lumley, J.L., 1972. A First Course in Turbulence. MIT Press, 300 pp.