### **Adaptive Modeling and Adaptive Sampling Research**

**For Coupled Air-Sea Predictions** 





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http://www.deas.harvard.edu/~pierrel

- 1. Adaptive Modeling (Monterey Bay and California Current System region)
- 2. Adaptive Sampling
- 3. PLUSNet Research during FAF-05: Coupled Acoustical-Physical Adaptive Sampling
- 4. Conclusions

NURC High-resolution Coupled Coastal Prediction Workshop November 29, 2005



Robinson A.R. and P.F.J. Lermusiaux (2002). DA for physical-biological interactions. The Sea, Vol.12.

#### Atmospheric fluxes from 3km and hourly COAMPS (J. Doyle, NRL): Winds

#### Sensitivity to horizontal resolution

3 km improves Representation of Coastal Jets & Coastal Shear Zone



Our evaluations: e.g. Buoy winds (blue) vs COAMPS 72h forecasts (red dots)



But: Wind-stress curl (for ocean upwelling)?, Long-wave radiations (cloud effects)?

### Surface Temperature: 7 August-23 August

Illustrates

- Daily cycle
  - •Night/day T sequence
  - •Daily variability of rim currents
- Onset and maintenance of first upwelling state (Aug 7-18)
- Relaxation (Aug 19-23)



Nowcest: 7 Aug 2003

# **Adaptive Modeling: Motivations and Concepts**

- •Atmospheric and oceanic dynamics can be intermittent and highly variable, and can involve interactions on multiple scales
- In general, fields and interactions that matter vary in time and space
- Model uncertainties can be (very) large (e.g. for biogeochemical processes)
- For efficient forecasting, model structures and parameters should evolve and respond quantitatively to new data injected into the running prediction system
  - Quantitative correction of model biases
  - Quantitative automated evolution of model structures as a function of model-data misfits
  - Quantitative comparison of competing models and better scientific understanding
  - Multi-model data assimilation
- Model quantity (parameters, structures, state-variables) said to be adaptive if its formulation, classically assumed constant, is made a function of data values
  - Physical regime transition (e.g., well-mixed to stratified) and evolving/unknown turbulent mixing parameterizations
  - Variations of biological assemblages with time and space (e.g., variable zooplankton dynamics, summer to fall phytoplankton populations, etc) and evolving biogeochemical rates and ratios

# Towards Real-time Physical Adaptive Models



- Different Types of Adaptation:
  - Physical model with multiple parameterizations evaluated in parallel (hypothesis testing)
  - Physical model with a single adaptive parameterization (adaptive evolution). Not sketched.
- Model selection based on quantitative dynamical/statistical study of data-model misfits
- Multi-model estimates: adaptive learning of errors of each model and combination based on maximum likelihood (examples carried out for SST of HOPS and ROMS)

### **Semi-Automated Real-time Physical Adaptation during AOSN2**

- Prior to AOSN2, PE model calibrated to four historical conditions likely to be similar to the unknown August 2003 conditions
- Ten days in the experiment: Forecasts a bit too geostrophic/too warm in upper-layers and larger-scale OBCs needed
- Real-time Adaptation
  - SBL mixing parameters and Open Boundary Conditions (OBCs) adapted to new 2003 data
  - 49 sets of parameter values and OBC formulations evaluated
  - Configuration with smallest Bias/RMSE and highest PCC at data points selected
  - Improved upper-layer fields of Temp., Salinity and currents



#### **Parameters/Parameterizations Selected for Possible Improvement/Adaptation**

- i. Initial condition parameters and simulation restart time
- ii. SBL: parameters in vertical mixing and dissipation of atmospheric fluxes
- iii. Horizontal viscosities
- iv. Formulations of OBCs:

#### **Adaptation Procedure**

$$\begin{split} &\frac{\partial T_i}{\partial t} \pm c_n \frac{\partial T_i}{\partial n} = -\beta_i \; \frac{(T_i - T_i^{\ell}(x, y, z, t))}{\tau_i(z)} \\ &\frac{\partial u_i}{\partial t} \pm c_n \frac{\partial u_i}{\partial n} = -\beta_i \; \frac{(u_i - u_i^{\ell}(x, y, z, t))}{\tau_i(z)} \end{split}$$

- i. Parameter/parameterizations modified one at a time, then in groups
- ii. In total, 49 simulations ran in parallel in real time (starting from Aug 5 or 7, with DA up to Aug 15 and forecasts beyond (for Aug. 16, 17 and 18)
- iii. Bias, RMSE and PCC estimates computed at data points (glider data)
- iv. Model chosen: the one with smallest weighted sum of Bias/RMSE/PCCs

#### Result

	IC/Re-start date	$ u_0 $	$K^e_v$	$E_k$	Boundary Relaxation				
Non-adapted	OA on Aug. 7	50.0	5.0	0.15	None				
	of Aug. 2-7 data								
Adapted	"	"	6.0	0.22	Bnd <i>i</i>	y/n	$ au_i^b$	$ au_i^s$	$h_i$
					West	1	1.5	4.5	50
					South	1	1.0	2.0	50
					East	0	<u>-</u>	-	-
					North	1	1.0	2.0	50

 $v_0$ : shear viscosity at zero local gradient Richardson number (cm^2/s)

 $K_v^{e}$ : eddy diffusion for tracers within the wind-mixing depth  $h^e$  (cm<sup>2</sup>/s)

 $E_k$ : Ekman depth factor

Table 1

Parameters and parameterization before and after adaptation. Units:  $\nu_0$  ( $cm^2/s$ ),  $K_v^e$  ( $cm^2/s$ ),  $E_k$  (non-dim),  $\tau_i^b$  and  $\tau_i^s$  (days), and  $h_i$  (m) (see also App. A.1).



#### Bias, RMSE and PCC estimates for un-adapted (blue) and adapted (green) real-time physical models



Fig. 2. Comparison between real-time un-adapted (blue lines with a plus at each data point) and real-time adapted (green lines with filled circles) physical ocean model. (a) Bias estimate (Model - OAed data) for temperature and salinity, as a function of depth (m) and time (day). (b) As (a), but for the Root-Mean-Square-Error (RMSE) estimate. (c) As (a), but for the mesoscale Pattern-Correlation-Coefficient estimate. Comparisons were made at 20 depths: 2, 5, 10, 20, 30, 40, 50, 60, 70, 80, 90, 100, 125, 150, 175, 200, 250, 300, 350 and 400 m (maximum data depth). However, only values for the first 100 m are shown for visibility.

Fig. 3. Differences between the objectively analyzed (OAed) temperature data and the real-time un-adapted (left column) and real-time adapted (right column) temperature forecasts at 30 m, for Aug 16 (top), 17 (middle) and 18 (bottom), 2003. Differences are shown only where the expected error standard deviation of the OAed data is less then 30%. The region plotted varies from day to day, following glider

Fig. 4. Real-time un-adapted (left column) and real-time adapted (right column) temperature forecasts at 30 m, for Aug 16 (top), 17 (middle) and 18 (bottom), 2003.



Non-adapted

Adapted

Non-adapted

# Towards Real-time Adaptive Coupled Models



- Different Types of Adaptive Couplings:
  - Adaptive physical model drives multiple biological models (biology hypothesis testing)
  - Adaptive physical model and adaptive biological model proceed in parallel, with some independent adaptation
- Ongoing and Future Numerical Implementation
  - For performance and scientific reasons, both modes are being implemented using message passing for parallel execution
  - Mixed language programming (using C function pointers and wrappers for functional choices)

## **Oceanic Adaptive Sampling: Multiple Facets**

Foci	<ul> <li>Optimal ocean science (Physics, Acoustics and/or Biology)</li> <li>Demonstration of adaptive sampling value, etc.</li> </ul>			
Objective Fields	<ul> <li>i. Maintain synoptic accuracy (e.g. upwelling, atmosocean boundary layer)</li> <li>ii. Minimize uncertainties (e.g. uncertain ocean estimates), or</li> <li>iii. Maximize the sampling of expected events (e.g. start of upwelling/ relaxation, dynamics of upwelling filament, small scales/model errors)</li> <li>Multidisciplinary or not</li> <li>Local, regional or global, etc.</li> </ul>			
Time and Space Scales	<ul> <li>i. Tactical scales (e.g. minutes-to-hours adaptation by each glider/AUV)</li> <li>ii. Strategic scales (e.g. hours-to-days adaptation for glider/AUV group/cluster)</li> <li>iii. Experiment scales</li> </ul>			
Assumptions	<ul> <li>Fixed or variable environment (w.r.t. asset speeds)</li> <li>Objective field depends on the predicted data values or not</li> <li>Operational, time and cost constraints, or not, etc.</li> </ul>			
Methods	Bayesian-based, Nonlinear programming, (Mixed)-integer programming, Simulated Annealing, Genetic algorithms, Neural networks, Fuzzy logics			

For each of the 5 categories, there are multiple choices (only a few listed here) Choices set the type of adaptive sampling research

### a. Adaptive sampling via ESSE

- Objective: Minimize predicted trace of full error covariance (T,S,U,V error std Dev).
- Scales: Strategic/Experiment (not tactical yet). Day to week.
- Assumptions: Small number of pre-selected tracks/regions (based on quick look on error forecast and constrained by operation)
- Problem solved: e.g. Compute today, the tracks/regions to sample tomorrow, that will most reduce uncertainties the day after tomorrow.
- Objective field changes during computation and is affected by data to-be-collected
- Model errors  ${\it Q}$  can account for coverage term

Dynamics:	$dx = M(x)dt + d\eta$	$\eta \sim N(0, \mathbf{Q})$
Measurement:	$y = H(x) + \varepsilon$	$\varepsilon \sim N(0, \mathbf{R})$

Non-lin. Err. Cov.:

 $dP/dt = <(x - \hat{x})(M(x) - M(\hat{x}))^T > + <(M(x) - M(\hat{x})(x - \hat{x})^T > +Q$ 

tr

**Metric or Cost function**: e.g. Find future H<sub>i</sub> and R<sub>i</sub> such that

$$\underset{Hi,Ri}{Min} tr(P(t_f)) \qquad or \qquad \underset{Hi,Ri}{Min} \int_{t_0}^{t_f} tr(P(t)) dt$$

#### Which sampling on Aug 26 optimally reduces uncertainties on Aug 27?



- Based on nonlinear error covariance evolution
- For every choice of adaptive strategy, an ensemble is computed



4 candidate tracks, overlaid on surface T fct for Aug 26



#### Best predicted relative error reduction: track 1



### **b. Optimal Paths Generation for a "fixed" objective field** (Namik K. Yilmaz, P. Lermusiaux and N. Patrikalakis)

- Objective: Minimize ESSE error standard deviation of temperature field
- Scales: Strategic/Tactical
- Assumptions
  - Speed of platforms >> time-rate of change of environment
  - Objective field fixed during the computation of the path and is not affected by new data
- Problem solved: assuming the error is like that now and will remain so for the next few hours, where do I send my gliders/AUVs?
- Method: Combinatorial optimization (Mixed-Integer Programming, using Xpress-MP code)
  - Objective field (error stand. dev.) represented as a piecewise-linear: solved exactly by MIP
  - Possible paths defined on discrete grid: set of possible path is thus finite (but large)
  - Constraints imposed on vehicle displacements dx, dy, dz for meaningful path

#### Example: Two and Three Vehicles, 2D objective field (3D examples also done)

Grey dots: starting points White dots: MIP optimal end points





#### c. Dynamics Objective Fields: Flux and/or Term-by-term Balances

• Physical model: Primitive-Equation (PDE, x, y, z, t: HOPS)

	Horiz. Mom.	$rac{D\mathbf{u_h}}{Dt} + f\mathbf{e}_3\wedge\mathbf{u}_h = -rac{1}{ ho_0} abla_hp_w$ -	$+   abla_h \cdot (A_h  abla_h \mathbf{u}_h) + rac{\partial  A_v  \partial \mathbf{u}_h / \partial z}{\partial z}$
	Vert. Mom.	$ ho g + rac{\partial p_w}{\partial z} = 0$	
	Thermal en.	$\frac{DT}{Dt} = \nabla_h \cdot (K_h \nabla_h T) + \frac{\partial K_v  \partial T}{\partial z}$	$T/\partial z$
٦	Cons. of salt	$\frac{DS}{Dt} = \nabla_h \cdot (K_h \nabla_h S) + \frac{\partial K_v  \partial S_h}{\partial z}$	$/\partial z$
	Cons. of mass	$\nabla \cdot \mathbf{u} = 0$	Heat Flux Balances: 4 fluxes normal to each side
	Eqn. of state	$\rho(\mathbf{r},z,t)=\rho(T,S,p_w)$	averaged over first upwelling period Mean Fluxes (W/m2) over: August 6, 2003 - 10:30:00pm -> August 13, 2003 - 4:30:00am GMT
			Mean T. alongshore adv. flux (+ towards y/pole) x 10 <sup>6</sup> Mean T. alongshore adv. flux (+ towards y/pole) x 10 <sup>7</sup>

-250

-300

-400 L -122.6

-<sup>3</sup>West se

-122.55

-122.5

Lon

-122.45

-122.4









### d. Dynamics Objective Fields: Lagrangian Coherent Structures and their Uncertainties for the Aug 26-29, 2003 Upwelling Period





See: Lermusiaux and Lekien, Aug. 2005, In press.

for "Dynamical System Methods in Fluid Dynamics", Oberwolfach, Germany.

### e. Dynamics Objective Field: M-S. Energy and Vorticity Analysis

Two-scale window decomposition in space and time of energy eqns: 11-27 August 2003



Transfer of KE from large-scale to meso-scale







• Center west of Pt. Sur: winds destabilize the ocean directly.

• Center near the Bay: winds enter the balance on the large-scale window and release energy to the meso-scale window during relaxation. X. San Liang





## Persistent Littoral Undersea Surveillance Network (PLUSNet)

Lead: Kuperman, Schmidt et al.

#### End-to-end System components

- Adaptive Tactical and Environmental Assessment and Predictions with distributed network of fixed and mobile sensors for improved DCL
- Coordination via network control architecture and covert communications
- System level concept demonstration in three years

#### Harvard Research Thrusts

- Multi-scale and non-hydrostatic nested ocean modeling
- Coupled physical-acoustical DA in real-time
- Acoustical-physical nonlinear adaptive sampling with ESSE and AREA



# **Physical-Acoustical Predictions and Adaptive Sampling**

P.F.J. Lermusiaux, D. Wang (MIT) P.J. Haley, Jr., W.G. Leslie, H. Schmidt et al.

Thanks to:

NURC: E. Coelho, E. Nacini, A. Cavanna, P. Ranelli

Cro. Met. Service: M. Tudor

HU: A. Robinson



### **FAF05** Goals and Accomplishments

- 1. Initiate and test the coupling of HOPS, ESSE (HU) and AREA (MIT)
- 2. Issue physical-acoustical adaptive sampling recommendations every day
  - Capture the vertical variability of the thermocline (due to fronts, eddies, internal waves, etc)
  - Minimize the corresponding uncertainties.

Adaptive sampling plans computed based on 1-to-2 days forecasts of physicalacoustical fields and uncertainties

### **Adaptive Sampling in Vertical Cross-Sections**

AUV-Track Base Lines - For - Specific Sound-speed Features



#### **Composite Base Lines**







**Internal Wave** 



Thermocline

# **High-Resolution Nested Ocean Modeling Domains**



		<b>Mini-HOPS</b>	Elba		
Resolution		<b>100m</b>	<b>300m</b>		
<b>C!</b> =0	$nx \times ny \times nz$	89×114×21	106×126×21		
Size	Extent	8.8×11.3 km	31.5×37.5 km		
Domain center		42.59°N, 10.14°E	42.63°N, 10.24°E		
Domain rotation		0°	0°		
Speed	dt = 50s	90 minutes/(model day)	120 minutes/(model day)		
	dt=300s	15 minutes/(model day)	20 minutes/(model day)		



# **Example of Results of Adaptive Yoyo Control (Jul 20-21)**



#### **Multiscale Dependences of Coupled Ocean-Atmosphere Processes**

15N

10N

5N

- On atmos. large-scale, ocean SST usually negatively correlated to surface winds
- On atmos. mesoscales (100-3000 km), " positively "
  - On atmos. sub-mescocales, ???



Fig. S5. Schematic illustration of wind variations in the marine atmospheric boundary layer (MABL) induced by small-scale features in the SST field (S2-S4). Cooling over cool water increases stratification and stabilizes the MABL, inhibiting vertical mixing by turbulence and convection. This decouples the surface winds from the winds aloft and increases the wind shear near the sea surface (blue curve). Heating over warm water decreases stratification and destabilizes the MABL, enhancing vertical turbulent mixing and convection. This deepens the MABL and mixes momentum downward from aloft, decreasing the wind shear near the sea surface (red curve). Deepening of the MABL over warm water has been observed directly (S5, S6) and inferred from satellite measurements of increased low-level cloudiness (S7-S9). Cooling and warming of the MABL also create an enhanced pressure gradient in the direction of the SST gradient (S10-S12). Small imbalances between the pressure gradient force and frictional effects from vertical turbulent mixing accelerate the flow in regions of strong downwind SST gradients, resulting in higher surface wind speeds over cool water.

D. Chelton et al, Science, 2004

(Chelton et al, 2004)



**Fig. 5.** Four-year averages (August 1999–July 2003) of the spatial high-pass–filtered curl of the wind stress (top) from Fig. 3 and of the SST and vector-average wind stress (bottom) for the eastern tropical North Pacific (left panels) and the western North Atlantic (right panels). The color scale for the wind stress curl is the same as in Figs. 1 to 3. The 4-year average SST fields were derived from satellite measurements by the TRMM (Tropical Rainfall Measuring Mission) Microwave Imager (TMI) (*45*). For clarity, the 25-km wind stress vectors are displayed with reduced resolution on a 1° by 1° grid. The white fringes along the continental margins are gaps in the satellite coverage (about 25 km for QuikSCAT and about 75 km for TMI) owing to land contamination of the microwave signals.

# **Conclusions for Coupled Air-Sea Predictions**

# Coupled Adaptive Sampling

- -Data sets dedicated to coupled modeling are needed
- -Both -comprehensive- data sets and -targeted- data sets for specific processes
- -Can be optimized with adaptive sampling

# • Coupled Adaptive Modeling

- -Hierarchy of modeling options need to be evaluated/tuned
- -From simple linear feedback to full fledged-models
- -Multiple types/scales of coupling: from waves to atmos. mesoscale/large-scale
- -Computational issues/research

# • Coupled atmospheric-oceanic-acoustic effects important

- Waves and sea surface
- Daily cycle can be very significant, including for coastal currents and hydrography
- Wind-curl most important for ageostrophic properties
- Long-wave radiation
- Impacts on multiple littoral fields: physics, biology, seabed